

LETTER • **OPEN ACCESS**

Long-term ENSO prediction with echo-state networks

To cite this article: Forough Hassanibesheli *et al* 2022 *Environ. Res.: Climate* 1 011002

View the [article online](#) for updates and enhancements.

You may also like

- [Monitoring the pendulum between El Niño and La Niña events](#)
Jingzhi Su, Tao Lian, Renhe Zhang et al.
- [Prediction of Indian rainfall during the summer monsoon season on the basis of links with equatorial Pacific and Indian Ocean climate indices](#)
Sajani Surendran, Sulochana Gadgil, P A Francis et al.
- [Causal effects of Indian Ocean Dipole on El Niño–Southern Oscillation during 1950–2014 based on high-resolution models and reanalysis data](#)
Thanh Le, Kyung-Ja Ha, Deg-Hyo Bae et al.

ENVIRONMENTAL RESEARCH CLIMATE



LETTER

Long-term ENSO prediction with echo-state networks

OPEN ACCESS

RECEIVED
25 March 2022

REVISED
22 June 2022

ACCEPTED FOR PUBLICATION
7 July 2022

PUBLISHED
21 July 2022

Forough Hassanibesheli^{1,2,*} , Jürgen Kurths^{1,2,3} and Niklas Boers^{1,4,5} 

¹ Research Domain IV—Complexity Science, Potsdam Institute for Climate Impact Research, Potsdam, Germany

² Department of Physics, Humboldt University Berlin, Berlin, Germany

³ Saratov State University, 83, Astrakhanskaya Str., 410012 Saratov, Russia

⁴ Earth System Modelling, School of Engineering & Design, Technical University of Munich, Munich, Germany

⁵ Department of Mathematics and Global Systems Institute, University of Exeter, Exeter, United Kingdom

* Author to whom any correspondence should be addressed.

E-mail: forough.hassanibesheli@pik-potsdam.de

Keywords: ENSO, artificial neural networks, echo state network

Original Content from
this work may be used
under the terms of the
[Creative Commons
Attribution 4.0 licence](https://creativecommons.org/licenses/by/4.0/).

Any further distribution
of this work must
maintain attribution to
the author(s) and the title
of the work, journal
citation and DOI.



Abstract

The El Niño–Southern Oscillation (ENSO) is a climate phenomenon that profoundly impacts weather patterns and extreme events worldwide. Here we develop a method based on a recurrent neural network, called echo state network (ESN), which can be trained efficiently to predict different ENSO indices despite their relatively high noise levels. To achieve this, we train the ESN model on the low-frequency variability of ENSO indices and estimate the potential future high-frequency variability from specific samples of its past history. Our method reveals the importance of cross-scale interactions in the mechanisms underlying ENSO and skilfully predicts its variability and especially El Niño events at lead times up to 21 months. This study considers forecasts skilful if the correlation coefficients are above 0.5. Our results show that the low-frequency component of ENSO carries substantial predictive power, which can be exploited by training our model on single scalar time series. The proposed machine learning method for data-driven modeling can be readily applied to other time series, e.g. finance and physiology. However, it should be noted that our approach cannot straightforwardly be turned into a real-time operational forecast because of the decomposition of the original time series into the slow and fast components using low-pass filter techniques.

1. Introduction

The El Niño–Southern Oscillation (ENSO) is the dominant variability mode of the global climate system at interannual time scales [1]. It is related to interannual temperature and pressure anomalies in the tropical Pacific Ocean. ENSO can be classified into three main variability phases, namely the warm (i.e. El Niño) and cool (i.e. La Niña) phases with sea-surface temperature (SST) anomalies substantially above and below average, respectively, as well as the remaining neutral phases [2]. The warm ENSO phases have typical yet irregular return periods between 3 and 7 years, rendering them challenging to predict [3]. In the neutral phase, trade winds blow east to west resulting in a pile-up of warm water masses at the western boundary of the tropical Pacific Ocean. The resulting east–west SST gradient causes air to ascend in the western Pacific and circulate back to the eastern boundary of the tropical Pacific Ocean, where it descends again. This atmospheric circulation system is called the Walker cell. During El Niño (La Niña) phases, this circulation is weakened (strengthened), leading to warm (cool) SST anomalies in the central and eastern parts of the tropical Pacific.

During the last decades, substantial progress has been made in assessing spatial patterns, the temporal dynamics, the underlying physical mechanisms, and possible changes of ENSO in response to global warming [4, 5]. Pronounced anomalies of ENSO, i.e. both El Niño and La Niña events, can trigger droughts, floods, heat waves, and other extreme weather conditions around the globe [6]. Skilful forecasts of ENSO variability provide crucial information for decision makers to reduce the negative socioeconomic impacts of strong ENSO anomalies. The predictability of ENSO at interseasonal and longer time scales has therefore

attracted substantial attention, using process-based general circulation models [7, 8] and statistical approaches [9, 10]. In recent years, also machine learning models have been demonstrated to provide skillful forecasts of ENSO variability [11–13]. Many linear statistical methods such as linear inverse models [14] and model analogs [15] have been developed to forecast ENSO variability. Since ENSO dynamics is arguably nonlinear [16], also nonlinear statistical models have been introduced [10, 17]. In addition to these data-driven inverse modeling approaches, several statistical methods, based on complex network theory, have been proposed to forecast El Niño events. It has been demonstrated that complex networks provide powerful tools for investigating the spatial co-variability patterns of climatic variables [18]. Regarding the prediction of different ENSO phases, this approach suggests that spatial co-variability in the tropical Pacific plays a crucial role in ENSO predictability. Nevertheless, as for the classical statistical approaches above, the forecast horizon of network-based approaches remains limited to 1 year [19–21].

Due to their distinctive ability to capture nonlinear relationships and their capability of handling large amounts of data, ANNs have recently been employed for a broad range of applications in the geosciences, including pattern recognition, classification, and signal prediction [22, 23]. Despite their need for very large training datasets, deep-learning models [24] have been employed to successfully classify El Niño events [25] and to predict ENSO variability [12, 13]. The most skillful ENSO prediction scheme to date, which outperforms process-based dynamical models at yearly time scales, is based on a convolutional neural network (CNN) architecture [12]. This model has high forecast skill (with correlation to the observed ENSO index above 0.5) for 17 months and thus outperforms operational forecasts using process-based models by 5 months. The need for large amounts of training data for such complex ANN models has been addressed by training on process-based climate model simulations of historical climate variability before training further on reanalysis data [12].

A specific kind of ANNs that has been developed for sequential data such as climate time series is recurrent neural networks (RNNs). Due to their recurrent nature, general RNNs are computationally expensive during training and are potentially prone to vanishing/exploding gradient problems. As a result, echo state networks (ESNs) [26, 27] have been proposed as simplified RNN models with substantially faster training processes in comparison to conventional RNNs. Due to the simplified training process of ESNs, they typically require smaller amounts of training data to learn the dynamics of the underlying complex system compared to other ANN architectures. ESNs have been applied successfully for forecasting deterministically chaotic dynamical systems [26, 28]. Here, we develop an ESN-based method to predict non-deterministic time series with comparably low signal-to-noise ratio, which frequently arises in geosciences.

The key assumption underlying our approach is that ENSO variability can be approximately decomposed into slow and fast modes, where the slow mode is only mildly perturbed by high-frequency forcing such as, e.g. westerly wind bursts or the Madden–Julian oscillation [29]. Several methodologies and techniques have been developed to decompose time series into slow and fast variability components, such as, e.g. singular spectrum analysis (SSA) [30, 31], spectral methods such as the Butterworth (BW) filter [32, 33], or moving average (MA) filters [34]. Upon a careful evaluation of their performance, we focus on the BW filter in the following. In our approach, we employ an ESN architecture to model the low-frequency component (LFC) of ENSO. To address the high-frequency components and perform a forecast of the full variability, we combine the ESN with the past noise forecasting (PNF) [17] method, which samples suitable high-frequency components from past observations to force the low-frequency dynamics.

2. Data and method

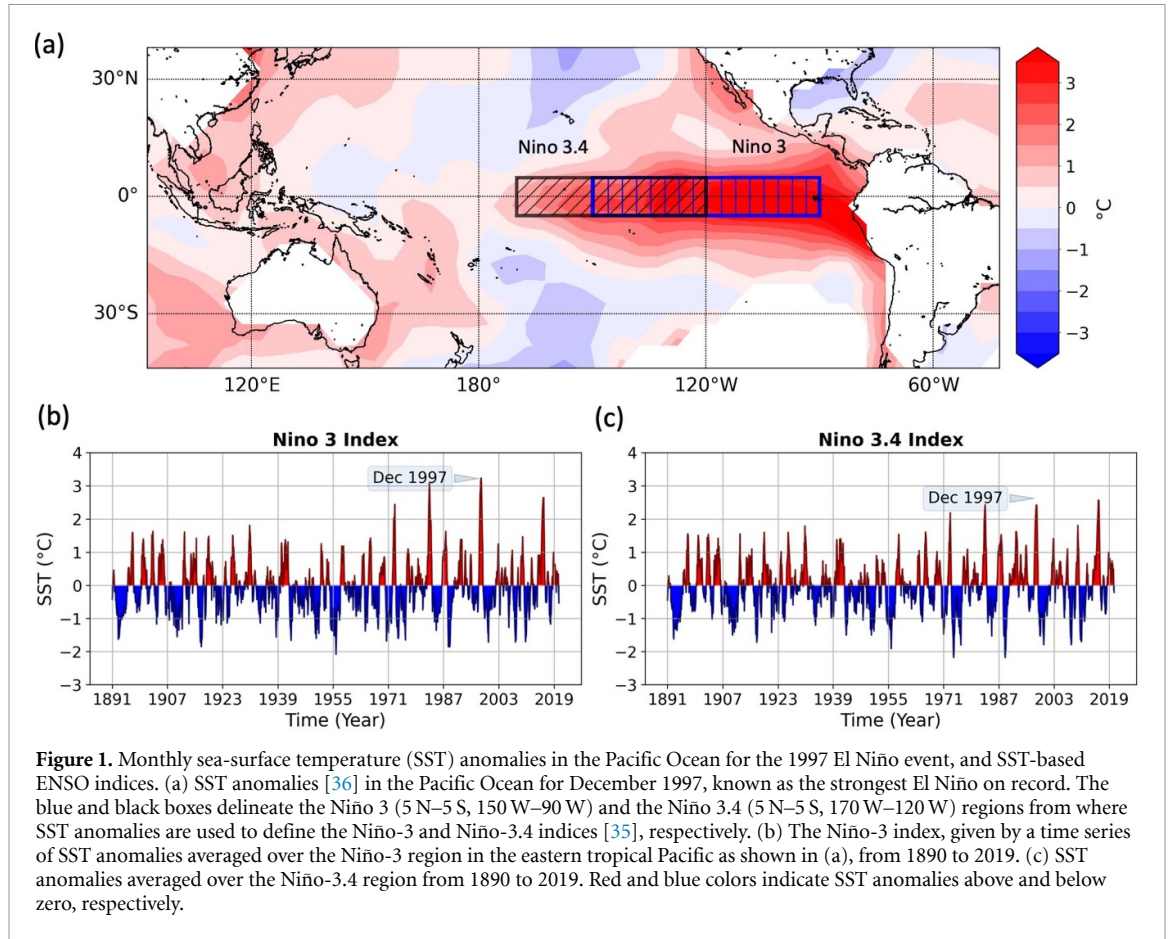
2.1. Data

We employ monthly SST anomalies for the time span 1891–2019 [35], averaged over the Niño-3 (5 N–5 S, 150 W–90 W) and Niño-3.4 (5 N–5 S, 170 W–120 W) regions in the tropical Pacific (figure 1(a)), which are commonly used to define the Niño-3 (figure 1(b)) and Niño-3.4 indices (figure 1(c)). To obtain the monthly indices, the seasonal cycles were removed by subtracting the 1981–2010 monthly means from the corresponding monthly values. The resulting time series, were then standardized.

2.2. Echo state networks (ESN)

A supervised machine learning method that is suited for temporal and sequential data processing is given by ESNs [27, 37], a specialized type of RNNs with simpler implementation and training process. ESNs are comprised of three parts: an input layer, a so-called reservoir that processes the input, and an output layer that uses the reservoir output for prediction. The reservoir can be considered as a high-dimensional dynamical system with states $r(t)$ that evolve according to the following equation:

$$r(t + \Delta t) = \mathcal{F}(W_r r(t) + W_{in} x(t)) . \quad (1)$$



Here $W_{in} \in \mathbb{R}^{N \times L}$ is the input matrix mapping the input of dimension L to the reservoir space of dimension N . $W_r \in \mathbb{R}^{N \times N}$ is a sparse, weighted adjacency matrix with elements drawn randomly from a normal distribution with zero means and \mathcal{F} refers to a nonlinear activation function for which we choose the hyperbolic tangent. The reservoir does not require remembering the entire temporal information of the past; hence, it skips unnecessary information after some time, resulting in extremely fast converging of the training. In this architecture, an input $x(t)$ of dimension L is non-linearly embedded in the reservoir through a randomly chosen $(N \times L)$ matrix W_{in} . Contrary to most conventional NN models, in the ESN setup, the values of W_{in} and W_r are not optimized but fixed at randomly chosen values. Hence, the only trainable connections are given by the reservoir-to-output layer matrix W_{out} , through which the reservoir states are mapped back to the L -dimensional outputs $y(t)$, i.e.

$$y(t + \Delta t) = (W_{out}r(t + \Delta t)) . \quad (2)$$

The goal of the system is to approximate the desired outputs x_d , i.e. to find optimal values for W_{out} that minimize the loss function \mathcal{L} . This optimization is most conveniently and robustly done using Ridge Regression (also known as Tikhonov Regularization) [38] with a regularization penalty parameter λ , which enhances the numerical stability and prevents over-fitting. The task is, thus, to minimize the cost function

$$\mathcal{L} = \sum_{t=0}^T ||W_{out}r(t) - x_d(t)|| + \lambda ||W_{out}||^2 . \quad (3)$$

Here $||..||$ is the L_2 -norm of a vector and the regularized optimal output matrix W_{out} is determined via:

$$W_{out} = (R^T R + \lambda \mathbb{I})^{-1} R^T x_d \quad (4)$$

where $R \in \mathbb{R}^{N \times L}$ and \mathbb{I} denotes the $N \times N$ identity matrix. The prediction phase shown in figure S5 (corresponding to $t \geq 0$) initiates with the information of the reservoir state $r(t^*)$ at the last time step of the training phase to make the prediction $y(t^* + 1)$. Further predictions are then iteratively made forward in time, by passing $y(t^* + 1)$ to the ESN as input to produce a forecast $y(t^* + 2)$, and so on. Specifically, we train the model between $[t^* - T, t^*]$ with a fixed training data length $T = 1092$ and then iteratively predict the

subsequent 25 months. This forecasting procedure is first carried out for t^* equal to January 1982, then repeated for February 1982, and so on until the end of the time series is reached.

As with all ANN models, various hyper-parameters of the ESNs have to be determined. The choice of the reservoir size (N), the sparsity (p), and the spectral radius (ρ) plays an important role for the efficiency of the learning process. In order to find optimal values for the hyper-parameters, we tested for different sets of parameters (N, p, ρ) (see B.1). Regarding the reservoir size N , it is generally expected that the larger the reservoir, the better the approximation of the underlying dynamics of the system, although this typically saturates. For our ESN we selected $N = 600$. Concerning sparsity, p determines the distribution of nonzero elements in the reservoir matrix W_r . In this study, we chose $p = 0.2$. Another hyper-parameter that can affect the ESN performance is the spectral radius ρ , that controls the echo state properties of the reservoir [37, 39, 40]. To set the spectral radius ρ of the reservoir matrix W_r , its entries are multiplied with the fraction of the desired spectral radius which is the largest absolute eigenvalue of the matrix W_r [39]. We choose $\rho = 0.95$. We employ these choices for the hyper-parameters consistently to initialise the model to learn the dynamics and predict the target SST time series.

2.3. Past noise forecasting (PNF)

Generally the pivotal problem of forecasting stochastic time series, such as the ones arising from measuring natural systems like climate, is to understand the dynamics of unresolved, high-frequency variables, since they may have considerable influence on the prediction of the resolved, low-frequency variables. A possible solution strategy is to estimate the time-dependent high-frequency forcing from suitable samples from its past history. Chekroun *et al* [17] originally developed a particular prediction methodology called PNF to circumvent this problem. The purpose of PNF is to select the best sample of past stochastic forcing to drive the system into the future, using the knowledge of the past noise trajectory. To find such potential noise realizations from past parts of the time series, the stochastic forcing is conditioned on the LFC of the underlying system. Here, we use BW low-pass filter to decompose the time series into an LFC and the corresponding high-frequency part. We then split the LFC into different segments of length Δ and search for those segments that resemble the reference LFC segment just preceding time step t^* , at which we initiate the forecast:

$$\text{RMSE}(\text{LFC}(t_j, t_j + \Delta) - \text{LFC}(t^* - \Delta, t^*)) \leq \alpha \quad (5)$$

$$\text{CORR}(\text{LFC}(t_j, t_j + \Delta) - \text{LFC}(t^* - \Delta, t^*)) \geq \gamma. \quad (6)$$

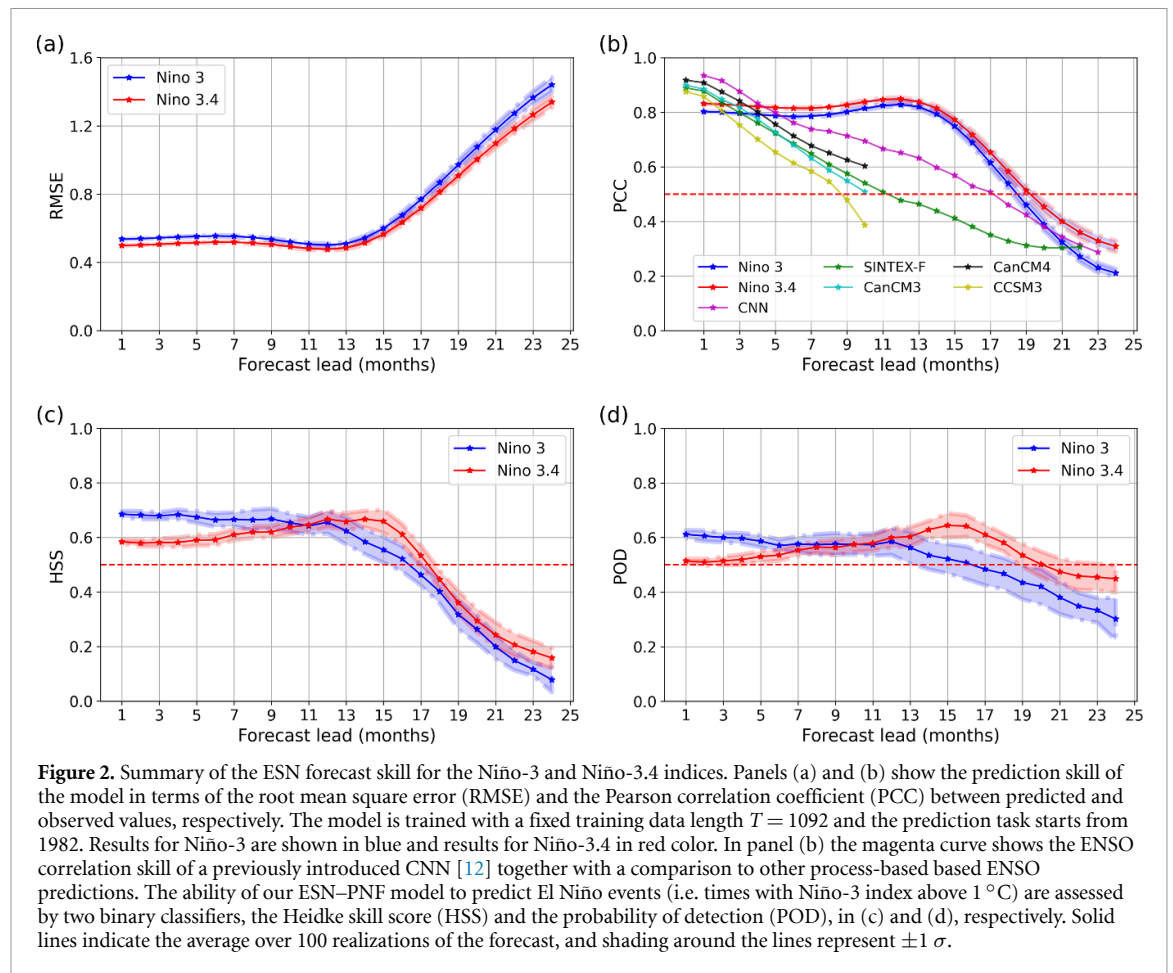
With these selection criteria, one can then identify high-frequency components starting at different $t_j < t^* - \Delta$, which we then use to force the ESN prediction for times $t > t^*$.

2.4. ESN implementation and combination with PNF

We train the ESN based on 80% of the original dataset (the training interval thus has a length of 1092 months) prior to t^* , the time that we aim to commence the prediction task (see figure S5). Using a BW low-pass filter, we decompose the training data into low- and high-frequency components, where the latter is treated as noise forcing. We then pass both components as input to the ESN. At each time, we perturb the LFC at time t by noise at time $t + 1$. During the training procedure, the weights of the output matrix W_{out} are optimized by minimizing the loss function using ridge regression [38]. Using the optimized W_{out} , the trained model starts recursively forecasting 25 consecutive months ahead of t^* .

In the prediction phase, since we do not have any information about the future behavior of the noise, we employ the past-noise forecasting (PNF) method to estimate past noise segments suitable to force the system into the future. As described in section 2.3, first we find those LFCs ($t_j, t_j + \Delta$) that resemble the LFC ($t^* - \Delta, t^*$) with $\Delta = 6$. We call the two LFC segments analogous if $\alpha \leq 0.5$ and $\gamma \geq 0.95$. In the next step, we select the corresponding high-frequency variability components of size $T = 25$. The prediction for the first time step is then, together with corresponding candidate noise analogues, passed to the ESN as inputs to predict the second time step, and so on. We repeat this training procedure for every month t_r starting from January 1891 with $r \in \{0, \dots, 432\}$ and predict 25 months ahead for each training set. After completing the prediction task for all training sets, we obtain the predicted time series at different lead times ranging from 1 to 25 months running from 1982 to 2019. For every single iteration, the predicted values are not seen by the model during the training phase.

Note that the skill of our ESN-based prediction scheme is slightly affected by varying the cutoff value C of the filter used to separate the low- from the high-frequency components of the ENSO index. Our results are



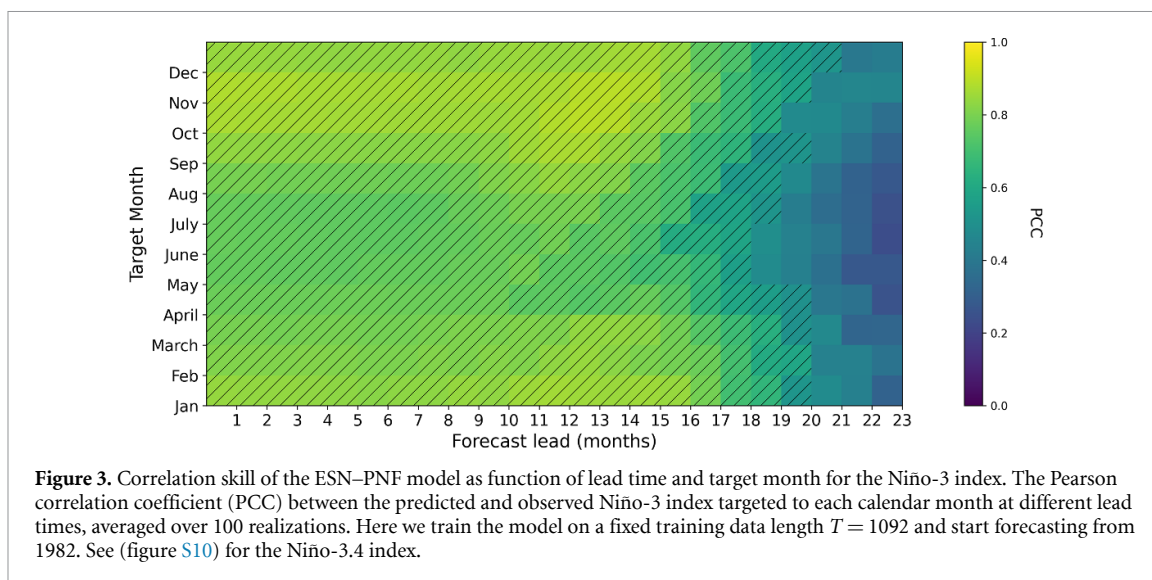
presented here for the optimal choice in terms of the overall forecast skill based on the above metrics ($C = 0.03$), and compared to alternative choices in (figure S13).

3. Results

Our ESN implementation consists of two main steps; the training phase during which a loss function is minimized to find optimal output weights for the ESN model, and the prediction phase during which the optimized ESN is used to predict unseen data (see figure S5).

We focus on forecasting the Niño-3 and Niño-3.4 (figure 2) indices during the time span from 1982 to 2019. To evaluate the skill of our ESN-PNF approach, we consider four different metrics, namely (I) the root mean square error (RMSE) and (II) the Pearson correlation coefficient (PCC) to evaluate the overall performance in predicting the ENSO index, as well as (III) the Heidke-skill score (HSS) and (IV) the probability of detection (POD) to evaluate the binary prediction of El Niño events (see A). For Niño-3 the RMSE remains almost constant at around 0.5°C for lead times between one and 14 months, and then increases linearly up to around 1.4°C at 24 months lead time (figure 2(a)). The PCC between the observed and simulated time series for the prediction phase remains constant around 0.8 also up to lead times of around 14 months and drops below 0.5 after 18 months (figure 2(b)). Note that, following [12] we consider forecasts as skillful if the correlation coefficients is above 0.5. To assess the skill of our model in predicting El Niño events (i.e. time steps with ENSO index larger than one), we applied the HSS (figure 2(c)) and the POD (figure 2(d)), cf appendix for details on these scores. Both event-based metrics also indicate high skill up to forecast lead times of 18 months. For the prediction phase starting at 1992, we obtain even a much higher skill during the first 14 months (figures S6 and S7). Our results remain similar when 5 month running averages of the Niño-3 and Niño-3.4 indices are considered, and when El Niño events are defined as time spans for which these smoothed indices are above 0.5°C for six consecutive months (see figure S8).

Note that the prediction score of our ESN-PNF model varies moderately over time, depending on the absolute values of the employed ENSO index. For example, we observe higher RMSEs during the 1997–98 and 2015–2016 El Niño events (figure S9). Hence, for El Niño and La Niña events with particularly high or



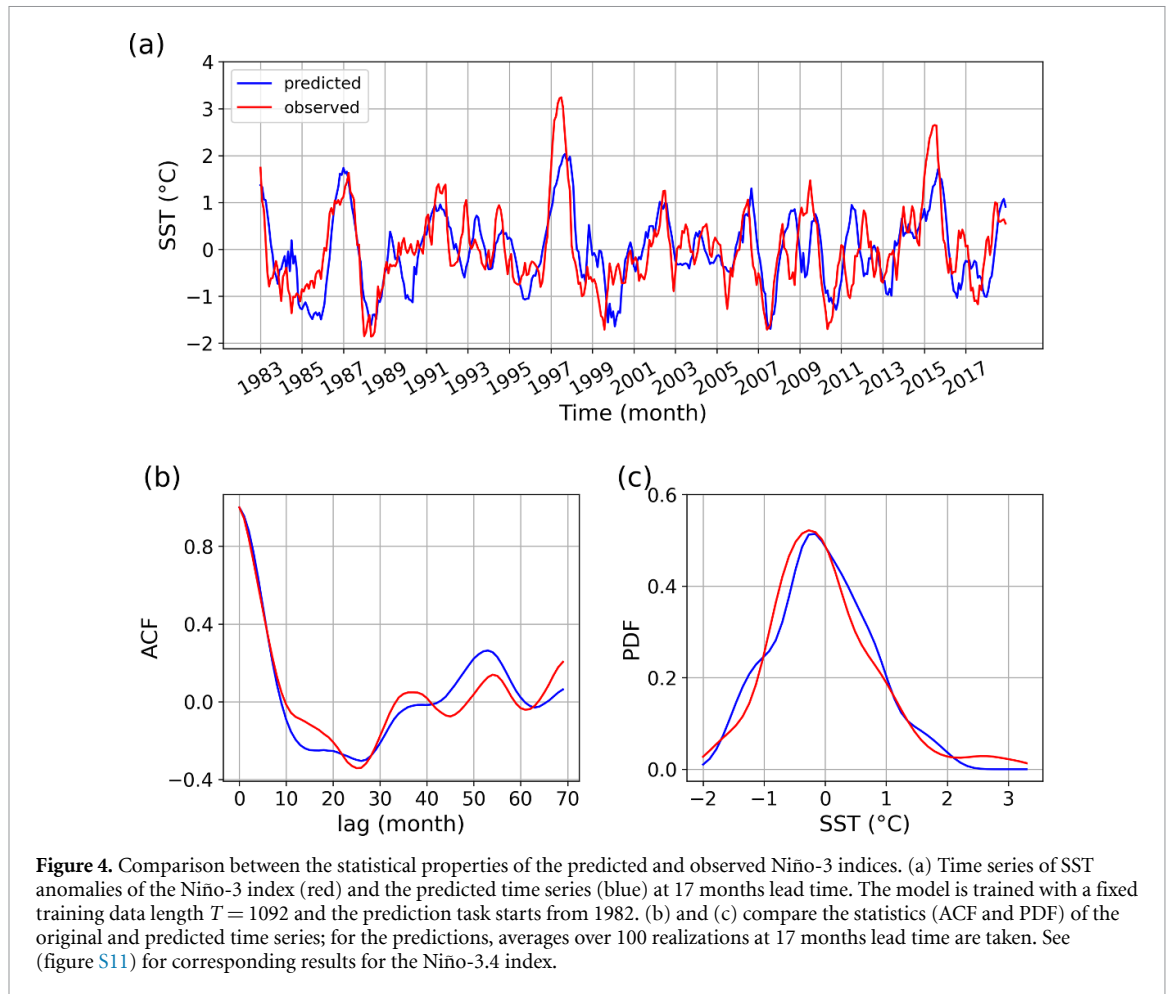
low values of the ENSO index, the RMSE is comparably higher, but the binary prediction of these events is still successful.

In order to further assess the reliability of our ENSO forecast, we investigate the sensitivity of the model's forecast skill to the target months at different lags (figures 3 and S10). In agreement with previous studies [12], our ESN–PNF model exhibits the longest forecast horizon for target months in boreal winter (the correlation between the observed and simulated ENSO index is above 0.5 for 21 months) and the shortest forecast horizon (18 months) for target months in late boreal spring. This spring predictability barrier [41] may be a result of the comparably weak Walker circulation and susceptibility of the coupled ocean–atmosphere system to the external forcing at that time of the year. Nevertheless, a recent CNN model has a valid forecast horizon (i.e. lead times at which the correlation remains above 0.5) that is substantially longer than for process-based dynamical models (11 months vs. 4 months) for the May–June–July season. For the May–June–July season, the valid forecast horizon of our ESN–PNF model is even substantially longer compared to the CNN model [12] (18 months vs 11 months). This indicates that our method is less affected by the spring predictability barrier compared to the CNN model and substantially less affected than process-based models, possibly because we estimate the future fast variables that drive the system using the PNF method.

As an example, the predicted SST anomaly between 1983 and 2019, at a lead time of 17 months, shows high visual resemblance to the observational Niño-3 index (figure 4(a)), although this is the longest lead time for which we consider our forecast to exhibit skill. The predicted values are computed as averages over 100 realizations at 17 month lead. The performance of our model is also examined by comparing the statistical properties such as the autocorrelation function (ACF, figure 4(b)) and probability density functions (PDF, figure 4(c)) of the predicted and observational time series. The ACF shows very close resemblance between the observed and simulated indices in terms of their correlation structure, and according to the Kolmogorov–Smirnov test, the hypothesis that the underlying distributions of the observed and simulated indices are identical cannot be rejected ($p = 0.35$). Corresponding results for the Niño-3.4 are very similar (figure S11).

4. Summary and discussion

There have been considerable advances in machine learning-based approaches to predict ENSO variability. One of the critical obstacles that most deep learning models encounter in predicting climate phenomena such as ENSO is the unavailability of sufficiently long observational time series, hindering training of these models purely on observations. In contrast to deep ANNs, the comparably simple ESNs can learn the dynamics of the underlying system from comparably small amounts of training data. Here, we expanded the ESN approach to predict time series with high noise levels, as they may arise from real-world measurements and experiments. We specifically applied our ESN model to predict the SST-based Niño-3 and Niño-3.4 indices. We decomposed the considered ENSO indices into dominant low- and high-frequency variability by using low-pass filter techniques and trained the ESN model on the slow mode of the system. To model the effect of the high-frequency forcing on the low-frequency variability, we estimated the potential future



high-frequency forcing of the system by relying on PNF method, which substantially enhanced the skill of our ESN–PNF model in predicting El Niño events. This improved predictability shows that interactions across multiple time scales indeed play a crucial role in generating the dynamics of ENSO. At the same time, the long valid time of our statistical prediction, with horizons up to 21 months, reveals that ENSO is in principle predictable for much longer lead times than previously thought, despite the fact that we have trained our ESN–PNF model only on single scalar ENSO indices.

Several studies demonstrated that using preprocessing techniques such as low-pass filtering can enhance forecasting skill of ANNs by providing high-quality inputs [42–44]. In practice, however, filtering methods such as SSA, BW filters, or MA-based methods incorporate information from future values to determine the LFC. As for all approaches involving filtering prior to the predictive modelling, our approach can therefore not directly be turned into a real-time forecasting of ENSO variability. Of course, the filtering can in principle be applied solely to the training part, but in that case well-known edge effects, i.e. errors in estimating the LFC arising at the beginning and end of the time series, would hinder a skillful prediction. We outline improved methods to reduce such errors when estimating low-frequency time series components as important subject for future research, which would also improve the skill of the method proposed here in an operational real-time forecasting setting.

While traditional modelling approaches are often challenged by cross-scale interactions, our approach in fact exploits the dependencies between low- and high-frequency variability for a more accurate modelling and statistical prediction. Our methods should therefore also prove valuable in other fields where cross-scale interactions are relevant, such as Neuroscience or Finance. Our approach indeed outperforms existing statistical, process-based, and deep-learning based approaches at lead times beyond 1 year, with the caveat that for the reasons explained above, we have not made our predictions operational. Nevertheless, our results reveal remarkable long-term predictability of ENSO far beyond the spring barrier [45]. For future studies, it might be of interest to combine process-based models with machine learning models such as the one presented here, to combine the strengths of both approaches.

Data availability statement

The data that support the findings of this study are available upon reasonable request from the authors.

Acknowledgments

We thank Mickael Checkroun for helpful discussions. This work has been financially supported by the German Academic Exchange Service (DAAD) under Project No. 57299294. N B acknowledges funding from the VolkswagenStiftung, the European Union's Horizon 2020 research and innovation programme under Grant Agreement No. 820970 and under the Marie Skłodowska-Curie Grant Agreement No. 956170, as well as from the Federal Ministry of Education and Research under Grant No. 01LS2001A.

Authors contribution

FH and NB conceived and designed the study. FH performed the numerical analysis and wrote the manuscript together with NB. All authors discussed results and contributed to editing the manuscript.

Appendix A. Scores for forecast validation

A.1. RMSE and correlation coefficient

To quantify the skill of our time series forecast, we use two well-established metrics, namely the root-mean-squared error (RMSE) and the Pearson's correlation coefficient (PCC). RMSE measures the average error magnitude with quadratic weight, which gives a quantitative evaluation of the closeness of the predicted variables x^p compared to the reference variables x^r (see figures 2(a), S9 and S13(a)).

$$\text{RMSE} = \sqrt{\frac{\sum_{i=1}^n (x_i^r - x_i^p)^2}{n}}. \quad (\text{S1})$$

PCC measures the linear relationship between the original data and corresponding forecasts. PCC can take values between -1 and 1 , where the positive and negative signs correspond to the direction of the linear relationship.

$$= \frac{\sum_{i=1}^n (x_i^r - \bar{x}_i^r) \cdot (x_i^p - \bar{x}_i^p)}{\sigma_{x^r} \cdot \sigma_{x^p}} \quad (\text{S2})$$

where σ_{x^r} and σ_{x^p} are the standard deviation of the reference and predicted variables, respectively.

A.2. Heidke skill score (HSS)

To verify the skill of our model for the binary forecast of El Niño events (i.e. months with ENSO index larger than 1°C), we employ the HSS [46], which is widely applied in binary classification problems. This metric is a measure of accuracy of a forecast with respect to a randomly generated forecast, adjusted to predictions that are correct by chance:

$$\text{HSS} = \frac{\text{TP} + \text{TN} - \text{CRF}}{N - \text{CRF}} \quad (\text{S3})$$

where TP and TN stand for true positives and true negatives, respectively, N is the total number of possible events, and CRF indicates the number of correct random forecasts, which can be calculated as follows:

$$\text{CRF} = \frac{(\text{TP} + \text{FN})(\text{TP} + \text{FP}) + (\text{TN} + \text{FN})(\text{TN} + \text{FP})}{N}. \quad (\text{S4})$$

Negative HSS values imply that the forecast skill of the model is worse than a random forecast, $\text{HSS} = 0$ indicates that the forecast is just as good as the random forecast, and $\text{HSS} = 1$ would indicate a perfect binary forecast. We also consider the probability of detection (POD) as a second metric to evaluate our prediction of El Niño events, which as defined as

$$\text{POD} = \frac{\text{TP}}{(\text{TP} + \text{FN})}. \quad (\text{S5})$$

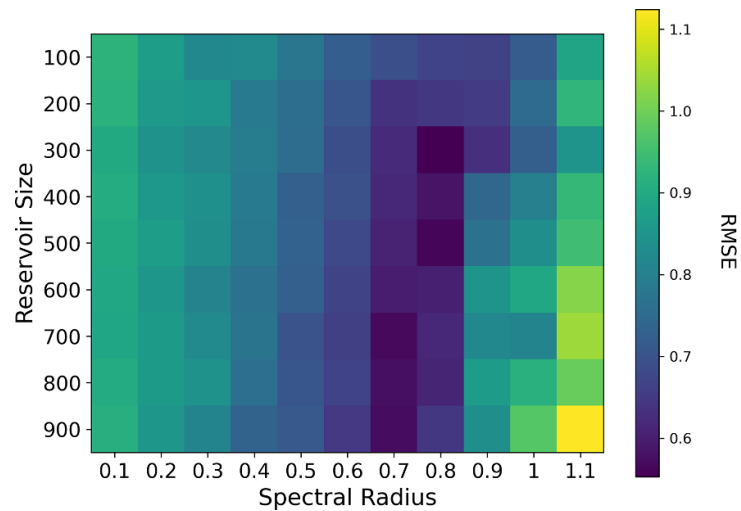


Figure S1. Echo state network performance with variations of reservoir size and spectral radius at fixed sparsity = 0.2.

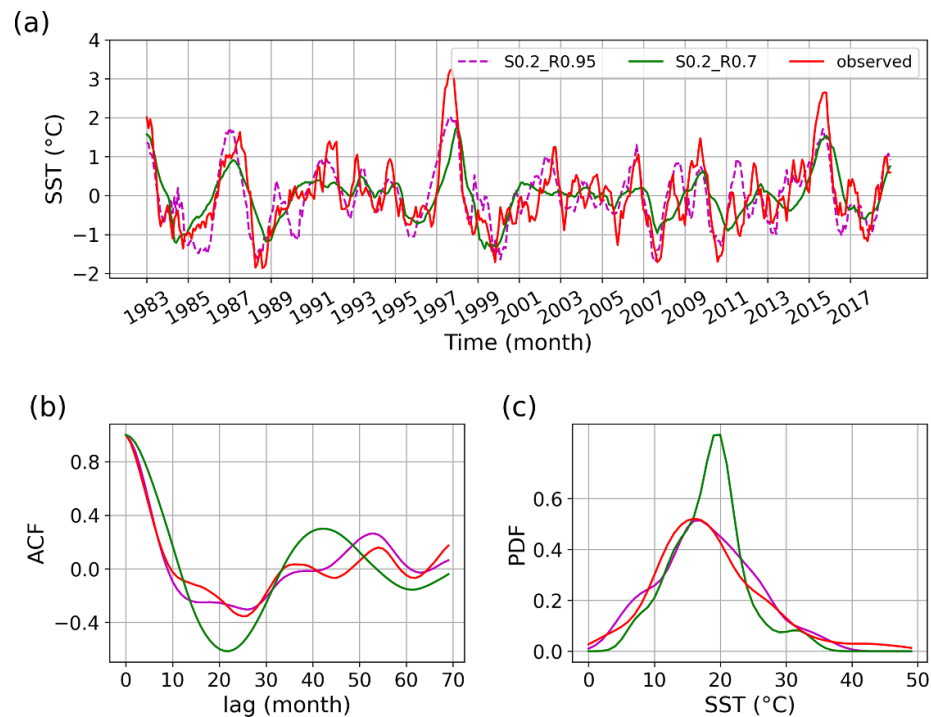


Figure S2. Comparison between the statistical properties of the predicted and observed Niño-3 indices. (a) Time series of SST anomalies of the Niño 3 index (red) and the predicted time series for ($\rho = 0.95$, sparsity = 0.2) (magenta), and ($\rho = 0.7$, sparsity = 0.2) (green), at 15 months lead time. (b) and (c) compare the statistics (ACF and PDF) of the original and predicted time series.

Appendix B. Supplementary figures

B.1. Hyper-parameters of the ESN–PNF

The key hyper-parameter that determines the echo state property of the ESN model is spectral radius (ρ). Figure S1 displays the performance of the ESN–PNF in terms of RMSE for different choices of reservoir size N and ρ values. Even though we observed better performance at $\rho = 0.7$ and $N = 600$, reducing the ρ value from 0.95 to 0.7 decreases the capability of the ESN model to produce the statistical properties of the underlying dynamics. To illustrate that, we compared the ACF and PDF of the predicted time series and original Niño-3 index at a 15 month lead time for $\rho = 0.95$ (magenta) and $\rho = 0.7$ (green) (see figure S2). According to our results, the reservoir scaled by $\rho = 0.95$ exhibits better performance in predicting the target signal.

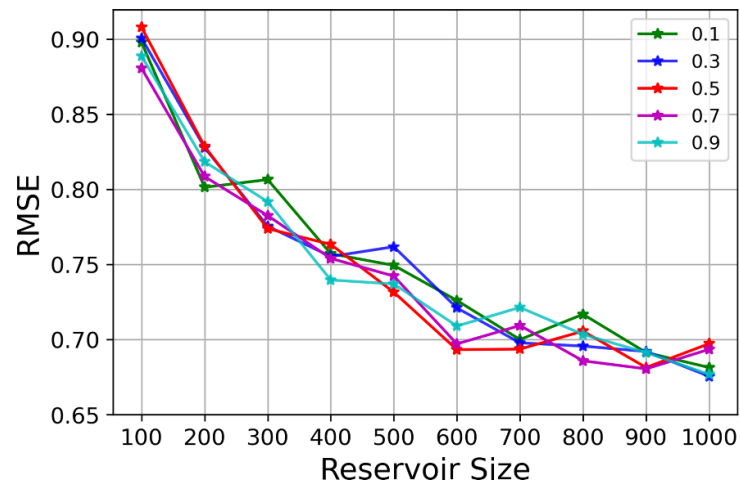


Figure S3. Echo state network performance with variations of reservoirs size with different sparsity.

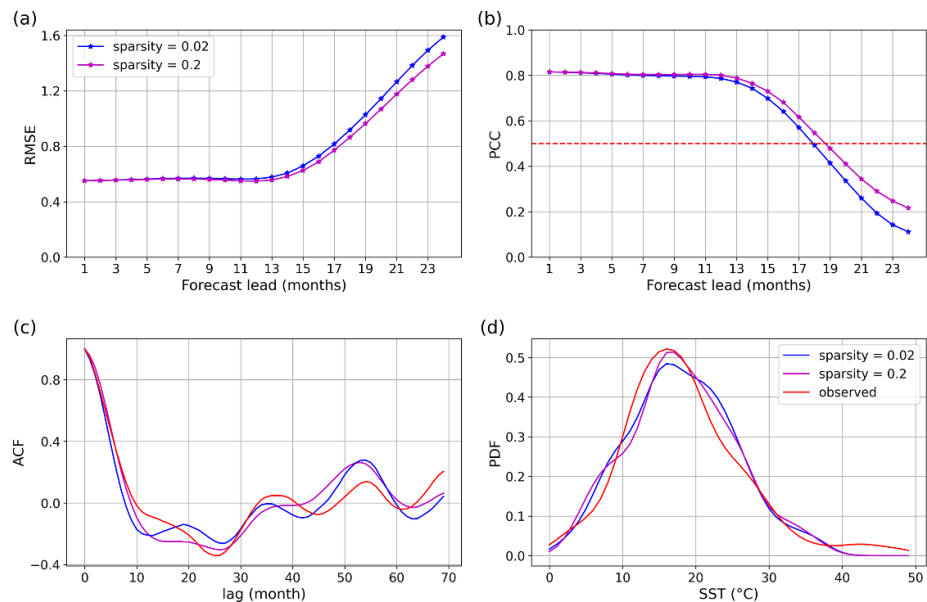
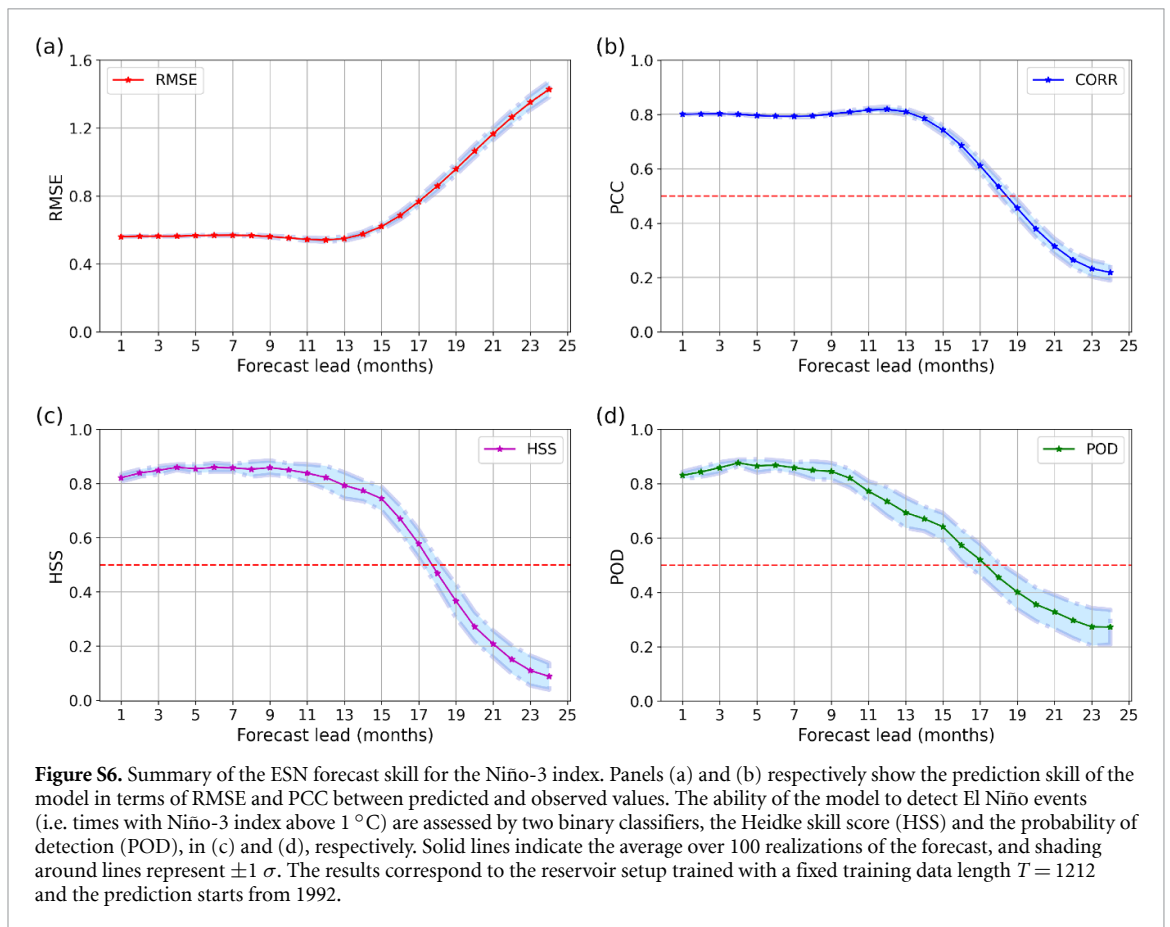
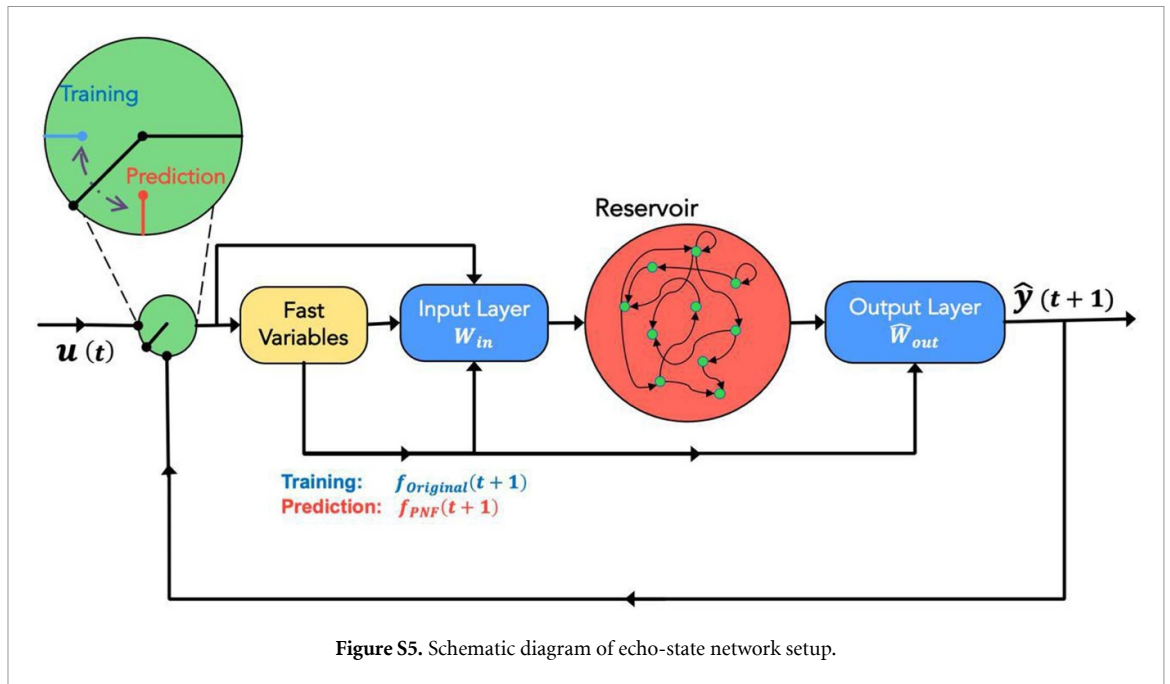


Figure S4. Comparison between the predicted and observed Niño-3 indices. Panels (a) and (b) show the prediction skill of the model in terms of the RMSE and the PCC between predicted and observed values for sparsity = 0.02 (blue), and sparsity = 0.2 (magenta) with $\rho = 0.95$. Panels (c) and (d) compare the statistics (ACF and PDF) of the original and predicted time series, respectively.

Further, we examined the role of reservoir size and connectivity on the model's performance (see figure S3) at the fixed $\rho = 0.95$. Our results demonstrate that increasing the reservoir size enhances the prediction performance of the model. We also observed that the model is not much sensitive to the sparsity of the reservoir. To have a better insight, we also compared the model's outputs with sparsity setting equal $p = 0.2$ and $p = 0.02$ in terms of different metrics (RMSE, PCC, ACF and PDF). As it can be observed from figures S4(a) and S4(b), the sparsity does not affect the performance of the ESN-PNF significantly.



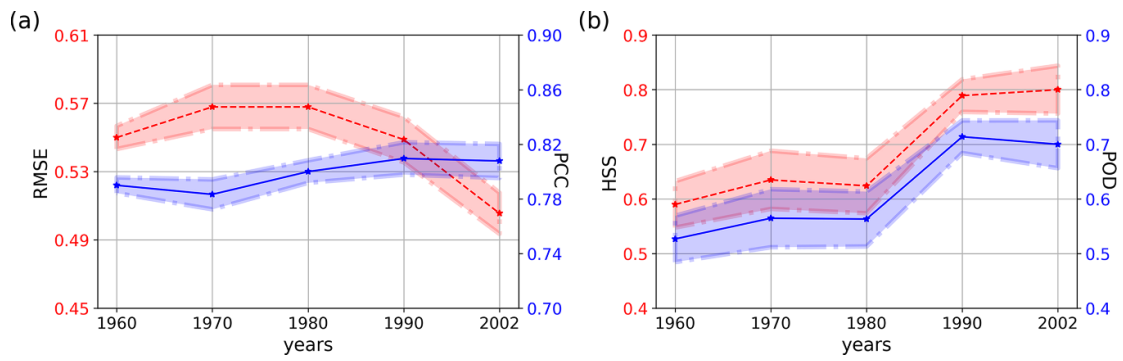


Figure S7. Dependency of the forecast skill on the training length. Panel (a) illustrates the RMSE and PCC values between the predicted and absolute values of Niño-3 index at a lead of 12 months with respect to different starting points. Panel (b) shows HSS and POD for the corresponding lead time. Our ESN model exhibits overall better performance the longer the training data. Shading around lines represents $\pm 1\sigma$.

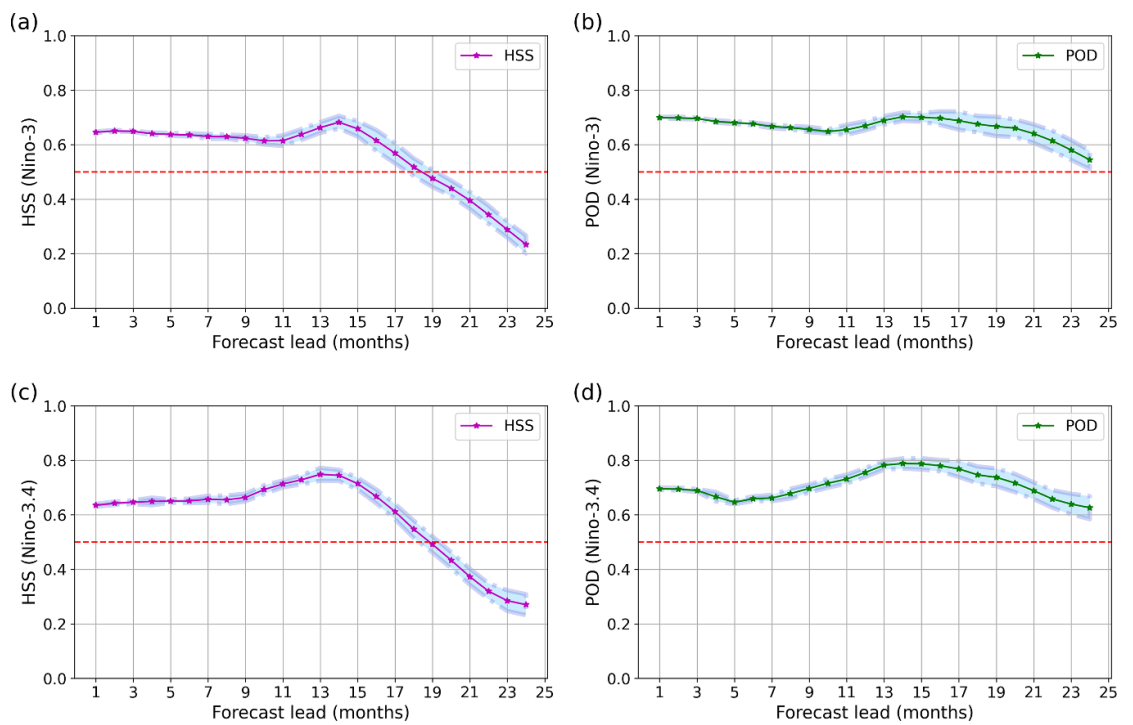
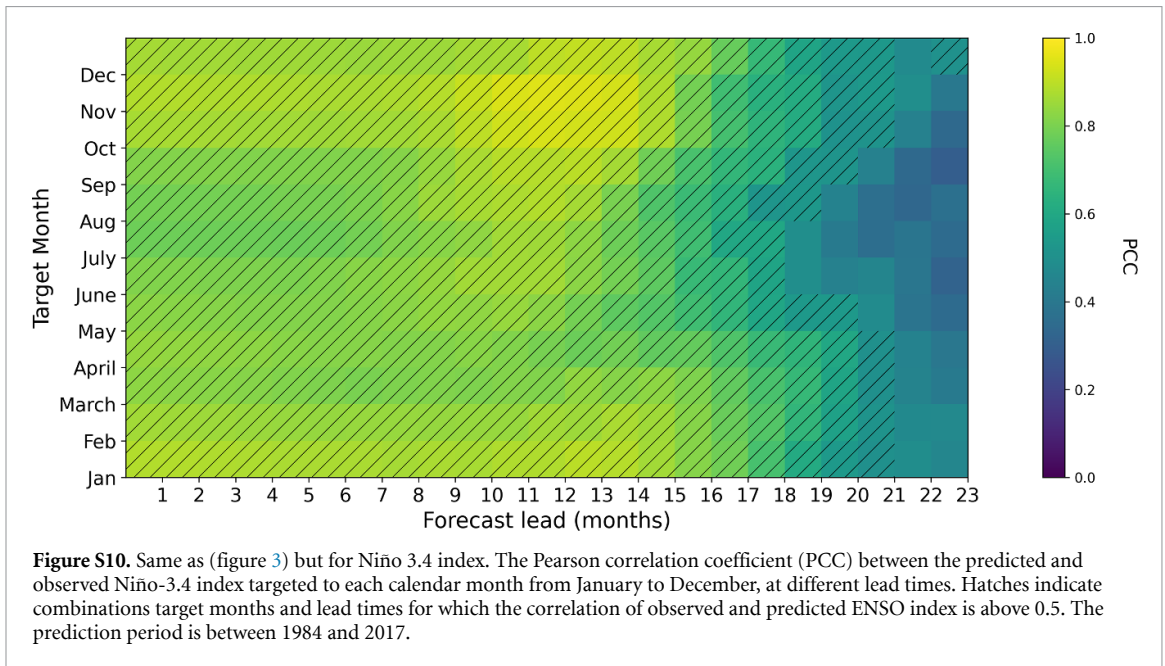
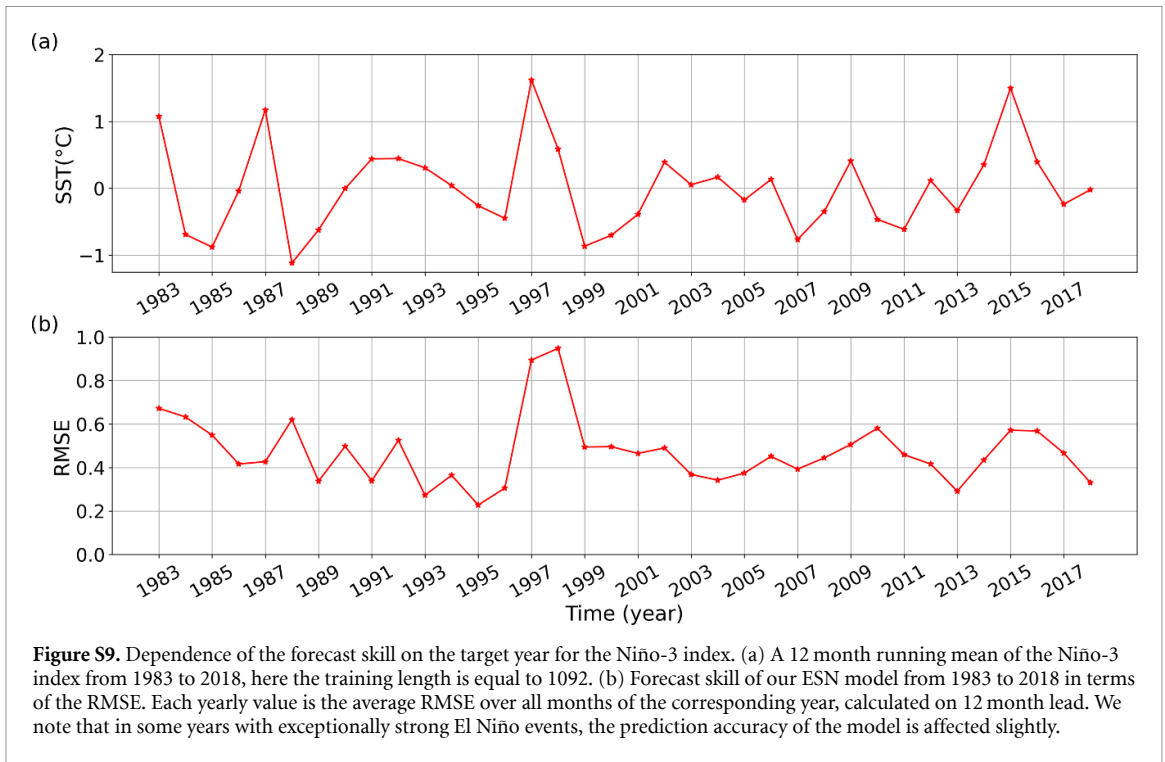
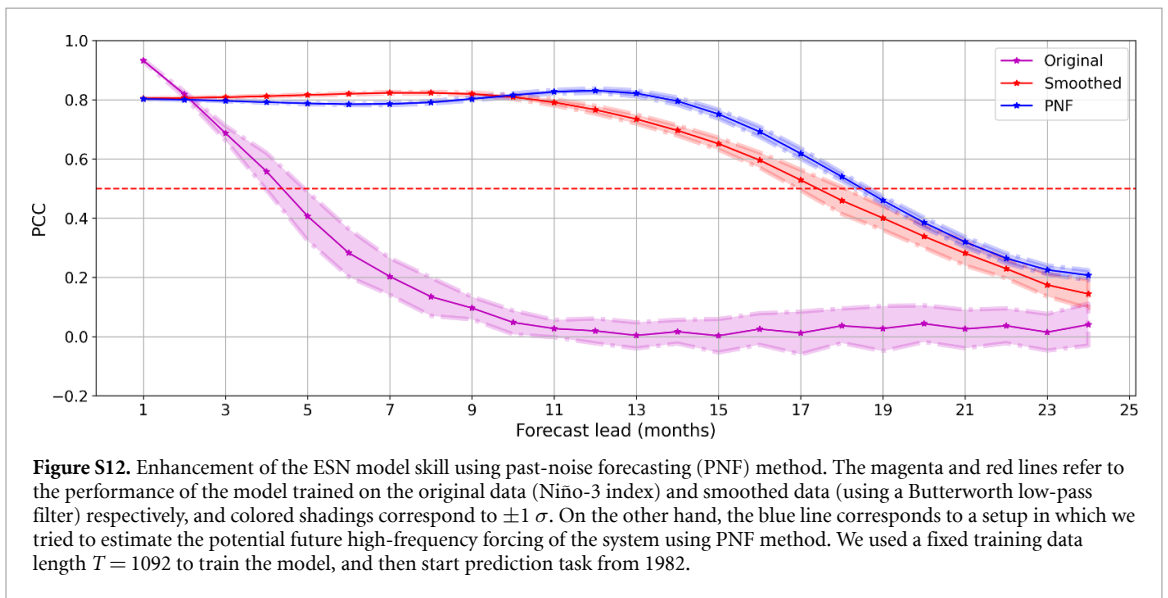
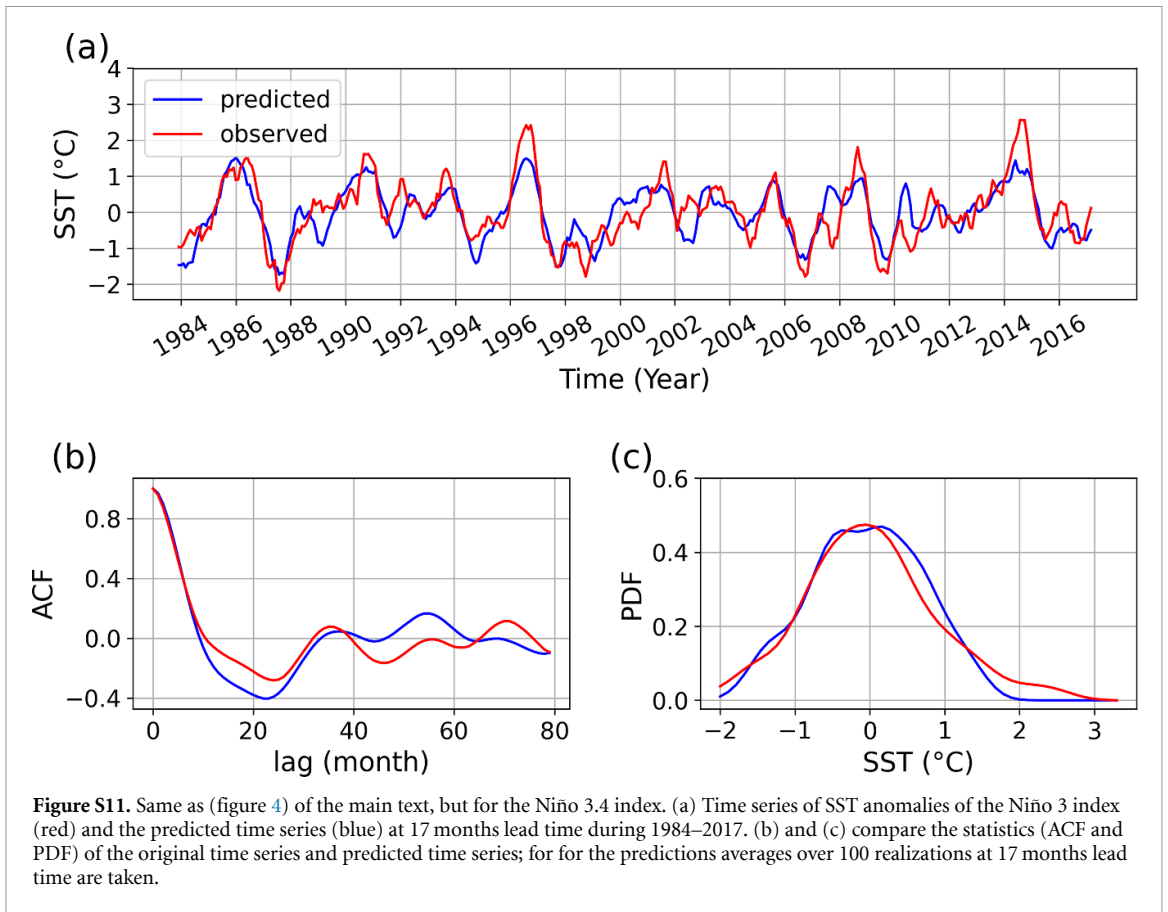
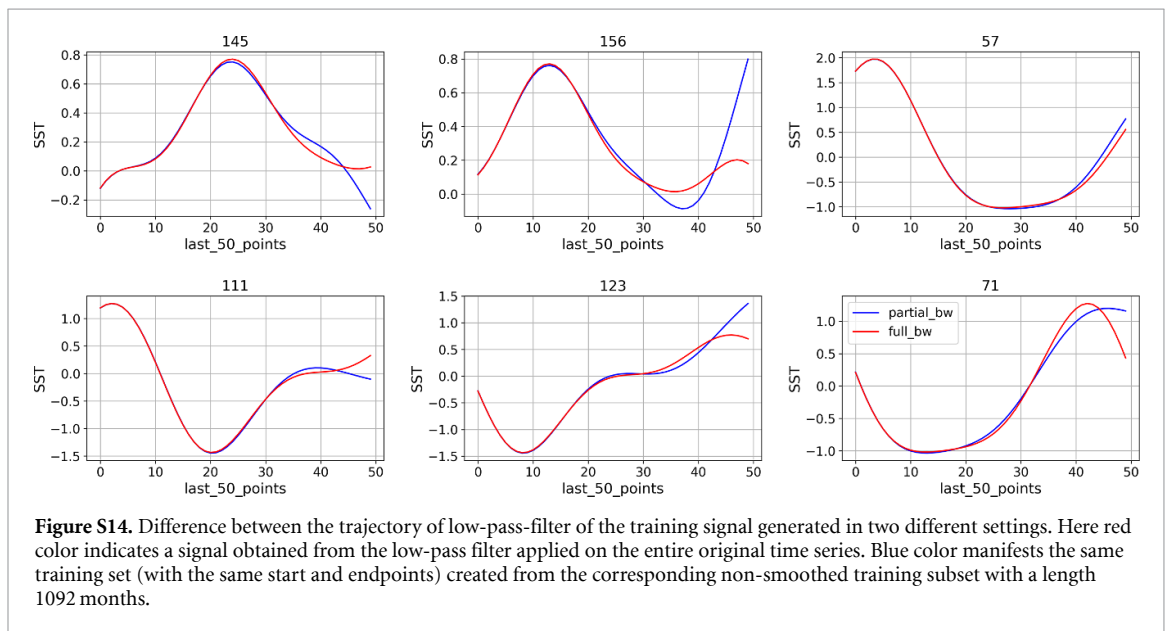
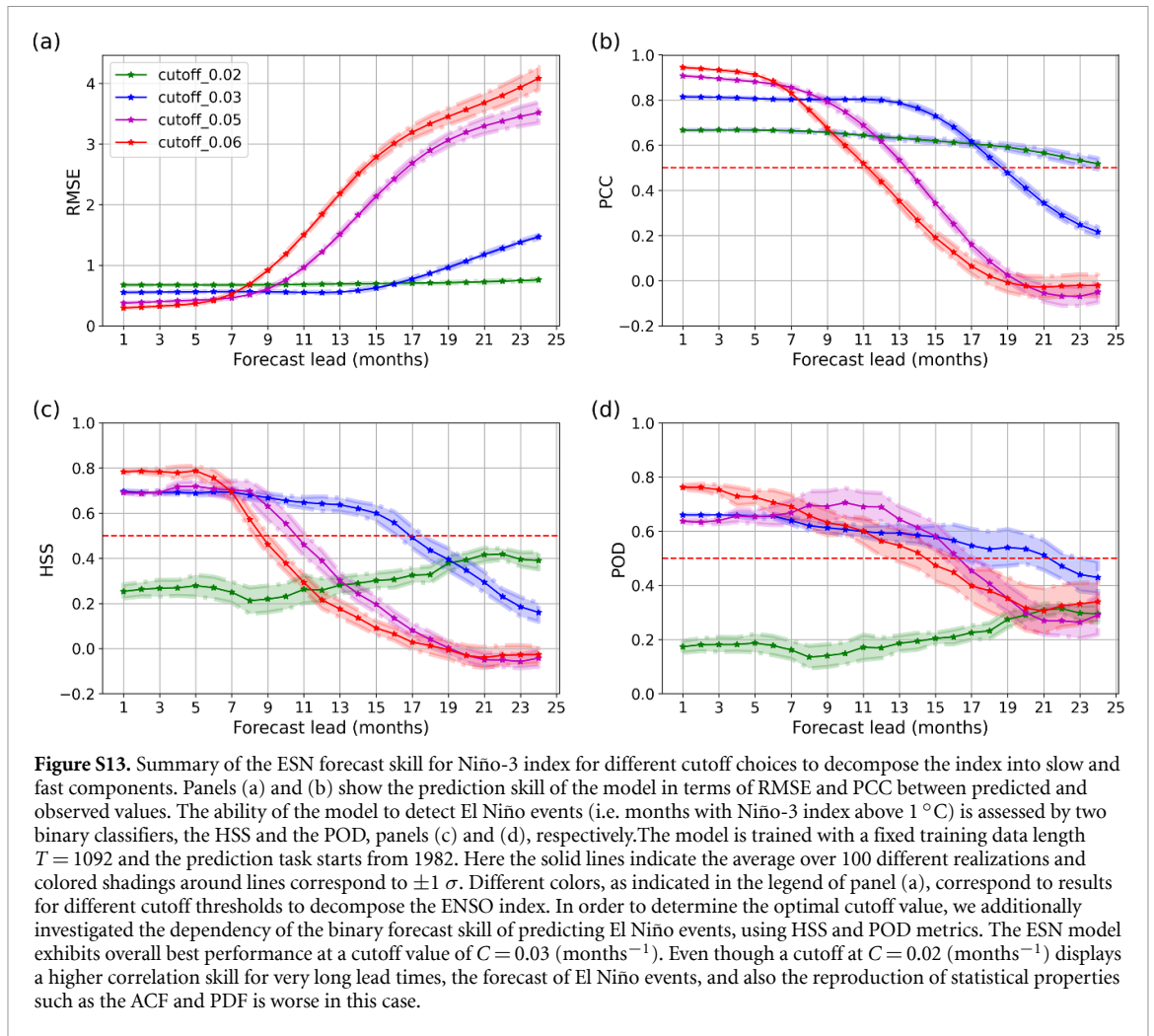
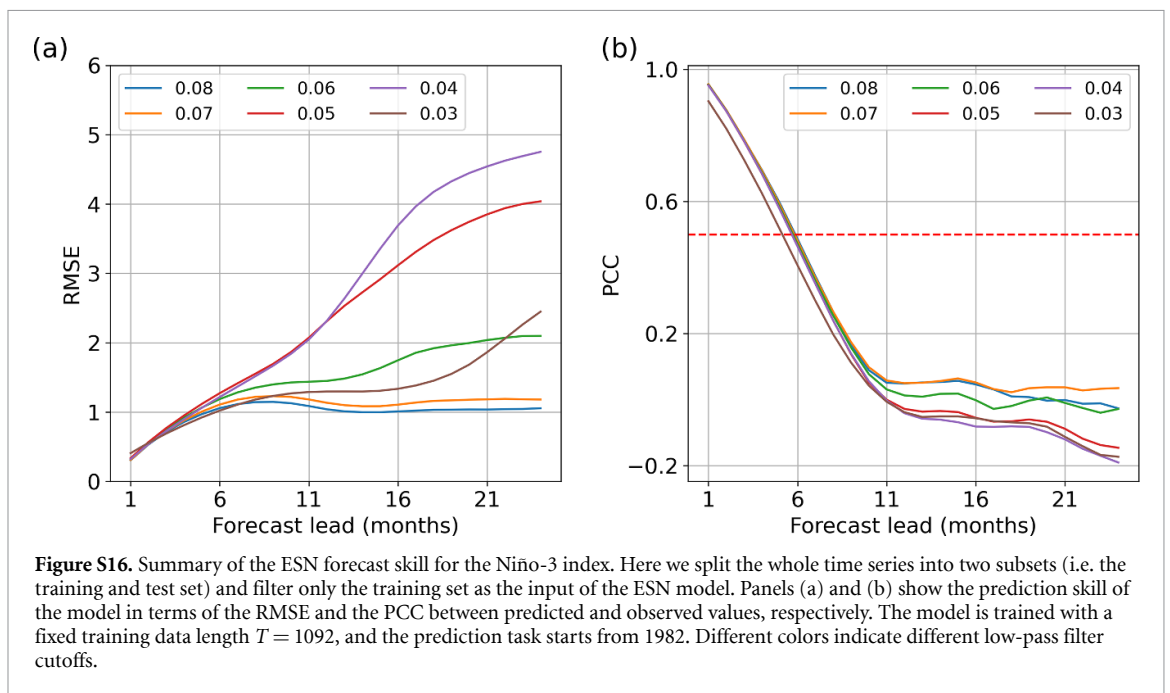
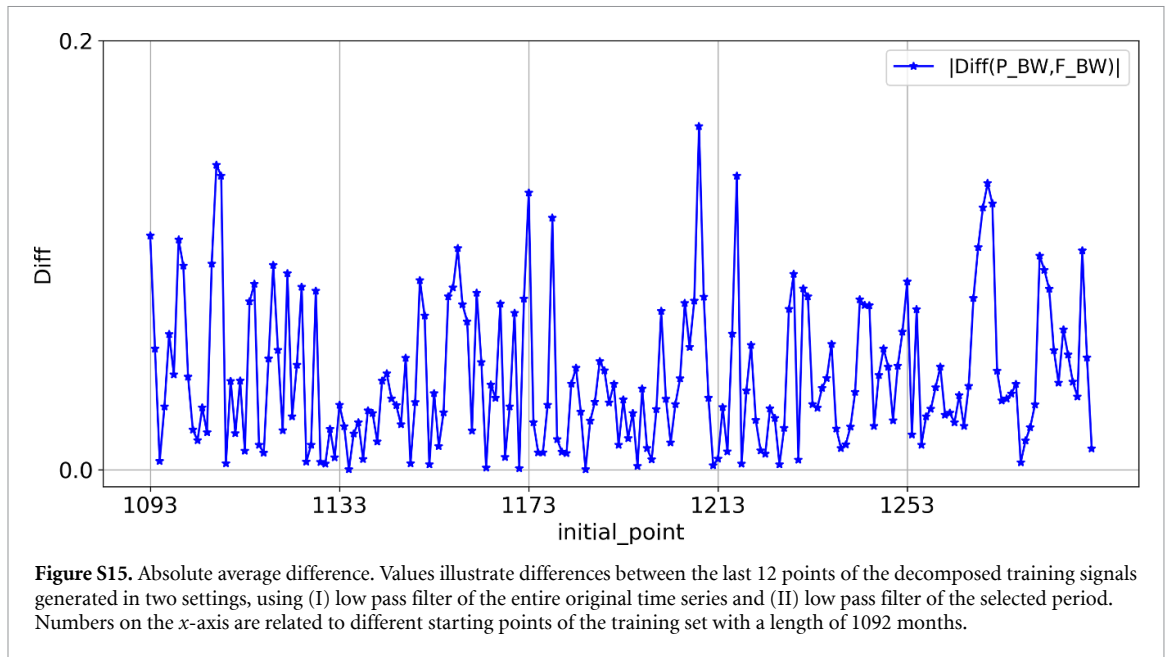


Figure S8. Detecting El Niño events for Niño-3 and Niño-3.4 indices with 5 month running mean average. The ability of the model to detect El Niño events (i.e., SST index above $0.5\text{ }^{\circ}\text{C}$) are assessed by two binary classifiers, the HSS and the POD. Panels (a) and (c) show HSS of the 5 month smoothed version of Niño-3 and Niño-3.4 indices respectively. The corresponding POD scores can be found in panels (b) and (d), where colored shadings around lines correspond to $\pm 1\sigma$.









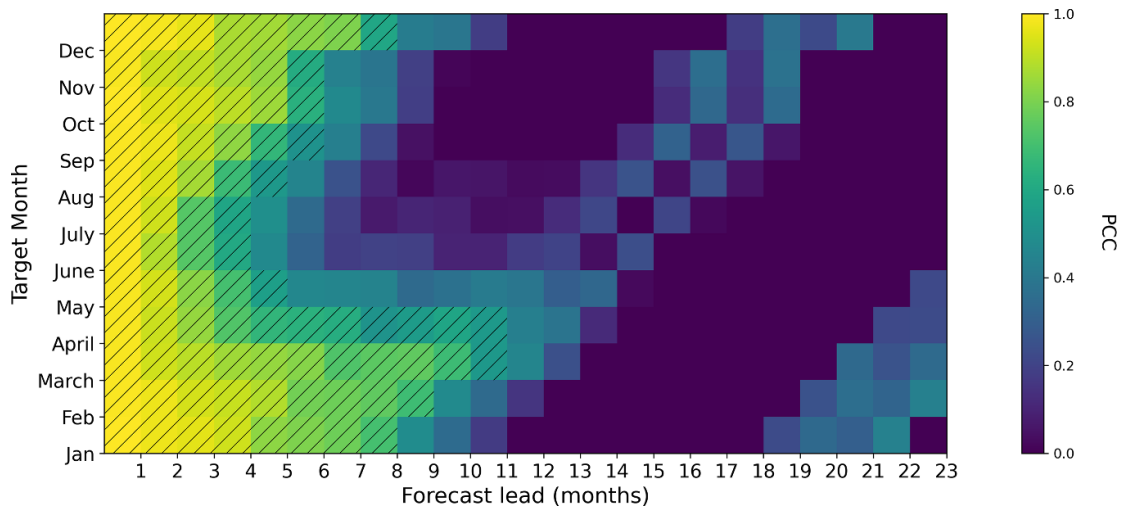


Figure S17. Correlation skill of the ESN trained on a subset preprocessed by low-pass filter, as a function of lead time and target month for the Niño-3 index. The PCC between the predicted and observed data targeted to each calendar month at different lead times. Here we train the model on a fixed training data length $T = 1092$ and start forecasting from 1982.

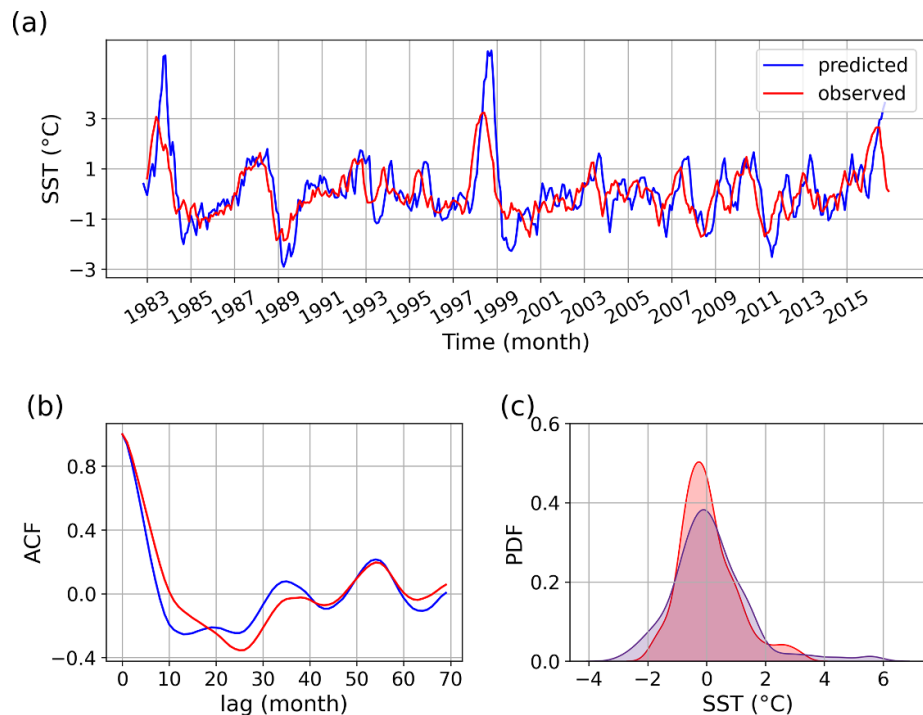


Figure S18. Comparison between the statistical properties of the predicted and observed Niño-3 indices. (a) Time series of SST anomalies of the Niño 3 index (red) and the predicted time series (blue) at 6 months lead time. The model is trained with a fixed training data length $T = 1092$ and the prediction task starts from 1982. (b) and (c) compare the statistics (ACF and PDF) of the original and predicted time series; for the predictions.

ORCID iDs

Forough Hassanibesheli  <https://orcid.org/0000-0001-6919-4358>

Niklas Boers  <https://orcid.org/0000-0002-1239-9034>

References

- [1] Philander S G H 1983 El Niño southern oscillation phenomena *Nature* **302** 295–301
- [2] Okumura Y M and Deser C 2010 Asymmetry in the duration of El Niño and la niña *J. Clim.* **23** 5826–43
- [3] Wallace J, Rasmusson E, Mitchell T, Kousky V, Sarachik E and Von Storch H 1998 On the structure and evolution of ENSO-related climate variability in the tropical pacific: lessons from toga *J. Geophys. Res.* **103** 14241–59

- [4] Latif M and Keenlyside N S 2009 El Niño/southern oscillation response to global warming *Proc. Natl Acad. Sci.* **106** 20578–83
- [5] Timmermann A *et al* 2018 El Niño–southern oscillation complexity *Nature* **559** 535–45
- [6] Lin J and Qian T 2019 A new picture of the global impacts of El Niño–southern oscillation *Sci. Rep.* **9** 1–7
- [7] Guilyardi E, Wittenberg A, Fedorov A, Collins M, Wang C, Capotondi A, Van Oldenborgh G J and Stockdale T 2009 Understanding El Niño in ocean–atmosphere general circulation models: Progress and challenges *Bull. Am. Meteorol. Soc.* **90** 325–40
- [8] Kirtman B P and Min D 2009 Multimodel ensemble ENSO prediction with CCSM and CFS *Mon. Weather Rev.* **137** 2908–30
- [9] Penland C and Magorian T 1993 Prediction of Niño 3 sea surface temperatures using linear inverse modeling *J. Clim.* **6** 1067–76
- [10] Kondrashov D, Kravtsov S, Robertson A W and Ghil M 2005 A hierarchy of data-based ENSO models *J. Clim.* **18** 4425–44
- [11] Dijkstra H, Hernandez-Garcia E, Lopez C and López C 2019 The application of machine learning techniques to improve El Niño prediction skill *Frontiers Phys.* **7** 153
- [12] Ham Y-G, Kim J-H and Luo J-J 2019 Deep learning for multi-year ENSO forecasts *Nature* **573** 568–72
- [13] Yan J, Mu L, Wang L, Ranjan R and Zomaya A Y 2020 temporal convolutional networks for the advance prediction of ENSO *Sci. Rep.* **10** 1–15
- [14] Penland C and Sardeshmukh P D 1995 The optimal growth of tropical sea surface temperature anomalies *J. Clim.* **8** 1999–2024
- [15] Ding H, Newman M, Alexander M A and Wittenberg A T 2019 Diagnosing secular variations in retrospective ENSO seasonal forecast skill using CMIP5 model-analogs *Geophys. Res. Lett.* **46** 1721–30
- [16] An S-I and Jin F-F 2004 Nonlinearity and asymmetry of ENSO *J. Clim.* **17** 2399–412
- [17] Chekroun M D, Kondrashov D and Ghil M 2011 Predicting stochastic systems by noise sampling and application to the El Niño–southern oscillation *Proc. Natl Acad. Sci.* **108** 11766–71
- [18] Boers N, Goswami B, Rheinwalt A, Bookhagen B, Hoskins B and Kurths J 2019 Complex networks reveal global pattern of extreme-rainfall teleconnections *Nature* **566** 373–7
- [19] Ludescher J, Gozolchiani A, Bogachev M I, Bunde A, Havlin S and Schellnhuber H J 2013 Improved El Niño forecasting by cooperativity detection *Proc. Natl Acad. Sci.* **110** 11742–5
- [20] Meng J, Fan J, Ashkenazy Y, Bunde A and Havlin S 2018 Forecasting the magnitude and onset of El Niño based on climate network *New J. Phys.* **20** 043036
- [21] Meng J, Fan J, Ludescher J, Agarwal A, Chen X, Bunde A, Kurths J and Schellnhuber H J 2020 Complexity-based approach for El Niño magnitude forecasting before the spring predictability barrier *Proc. Natl Acad. Sci.* **117** 177–83
- [22] Rasp S, Pritchard M S and Gentine P 2018 Deep learning to represent subgrid processes in climate models *Proc. Natl Acad. Sci.* **115** 9684–9
- [23] López E, Valle C, Allende H, Gil E and Madsen H 2018 Wind power forecasting based on echo state networks and long short-term memory *Energies* **11** 526
- [24] LeCun Y, Bengio Y and Hinton G 2015 Deep learning *Nature* **521** 436–44
- [25] Toms B A, Barnes E A and Ebert-Uphoff I 2020 Physically interpretable neural networks for the geosciences: applications to earth system variability *J. Adv. Model. Earth Syst.* **12** e2019MS002002
- [26] Jaeger H and Haas H 2004 Harnessing nonlinearity: predicting chaotic systems and saving energy in wireless communication *Science* **304** 78–80
- [27] Jaeger H 2007 Echo state network *Scholarpedia* **2** 2330
- [28] Pathak J, Hunt B, Girvan M, Lu Z and Ott E 2018 Model-free prediction of large spatiotemporally chaotic systems from data: a reservoir computing approach *Phys. Rev. Lett.* **120** 024102
- [29] Thual S, Majda A J, Chen N and Stechmann S N 2016 Simple stochastic model for El Niño with westerly wind bursts *Proc. Natl Acad. Sci.* **113** 10245–50
- [30] Vautard R, Yiou P and Ghil M 1992 Singular-spectrum analysis: a toolkit for short, noisy chaotic signals *Physica D* **58** 95–126
- [31] Hassani H 2007 Singular spectrum analysis: methodology and comparison *Journal of Data Science* **5** 239–57
- [32] Roberts J and Roberts T D 1978 Use of the Butterworth low-pass filter for oceanographic data *J. Geophys. Res.* **83** 5510–4
- [33] Gomez V 2001 The use of Butterworth filters for trend and cycle estimation in economic time series *J. Bus. Econ. Stat.* **19** 365–73
- [34] Alvarez-Ramirez J, Rodriguez E and Echeverría J C 2005 Detrending fluctuation analysis based on moving average filtering *Physica A* **354** 199–219
- [35] GCOS-AOPC/OOPC 2019 Working group on surface pressure (available at: https://psl.noaa.gov/gcos_wgsp/Timeseries/Nino3/) (Accessed 15 December 2019)
- [36] Kaplan A, Cane M A, Kushnir Y, Clement A C, Blumenthal M B and Rajagopalan B 1998 Analyses of global sea surface temperature 1856–1991 *J. Geophys. Res.* **103** 18567–89
- [37] Lukoševičius M and Jaeger H 2009 Reservoir computing approaches to recurrent neural network training *Comput. Sci. Rev.* **3** 127–49
- [38] Hoerl A E and Kennard R W 1970 Ridge regression: applications to nonorthogonal problems *Technometrics* **12** 69–82
- [39] Jaeger H, Lukoševičius M, Popovici D and Siewert U 2007 Optimization and applications of echo state networks with leaky-integrator neurons *Neural Netw.* **20** 335–52
- [40] Lukoševičius M 2012 A practical guide to applying echo state networks *Neural Networks: Tricks of the Trade* (Berlin: Springer) pp 659–86
- [41] Webster P J and Yang S 1992 Monsoon and ENSO: Selectively interactive systems *Q. J. R. Meteorol. Soc.* **118** 877–926
- [42] Wu C, Chau K W and Li Y S 2009 Predicting monthly streamflow using data-driven models coupled with data-preprocessing techniques *Water Resour. Res.* **45**
- [43] Kalteh A M 2017 Enhanced monthly precipitation forecasting using artificial neural network and singular spectrum analysis conjunction models *INAE Lett.* **2** 73–81
- [44] Wu C, Chau K W and Fan C 2010 Prediction of rainfall time series using modular artificial neural networks coupled with data-preprocessing techniques *J. Hydrol.* **389** 146–67
- [45] Lopez H and Kirtman B P 2014 WWBS, ENSO predictability, the spring barrier and extreme events *J. Geophys. Res.: Atmos.* **119** 10–114
- [46] Barnston A G 1992 Correspondence among the correlation, RMSE and Heidke forecast verification measures; refinement of the Heidke score *Weather Forecast.* **7** 699–709