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1 Targeted rainfall enhancement as an objective of forestation

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13
14 **Forestation efforts are accelerating across the globe in the fight against global climate**
15 **change, in order to restore biodiversity, and to improve local livelihoods. Yet, so far the**
16 **non-local effects of forestation on rainfall have largely remained a blind spot. Here we**
17 **build upon emerging work to propose that targeted rainfall enhancement may also be**
18 **considered in the prioritization of forestation. We show that the tools to achieve this are**
19 **rapidly becoming available, but we also identify drawbacks and discuss which further**
20 **developments are still needed to realize robust assessments of the rainfall effects of**
21 **forestation in the face of climate change. Forestation programs may then mitigate not**
22 **only global climate change itself, but also its adverse effects in the form of drying.**

24 Introduction

25 Forestation, including both afforestation and reforestation, is on the rise. The current decade is
26 the UN Decade on Ecosystem Restoration and to keep global warming ‘well below 2 degrees
27 Celsius’, as agreed in the Paris Agreement, large-scale forestation will likely be needed to
28 realize net negative emissions later this century (IPCC, 2022b). However, forestation is not a
29 silver bullet approach to halt climate change, and it is sometimes criticized for being ineffective
30 or even harmful (e.g. Calder, 2007; Veldman et al., 2015b). Researchers and practitioners are
31 making progress in understanding how, where and why forestation can be most effective and
32 should therefore be prioritized (e.g. Griscom et al., 2017; Brancalion et al., 2019; Hua et al.,
33 2022; Mo et al., 2023). Several recent studies developed optimizations of not only forest

34 conservation but also reforestation, considering different effects such as on biodiversity (Jung
35 et al., 2021; Strassburg et al., 2020). Further, ‘ten golden rules’ (Di Sacco et al., 2021) and
36 ‘fifteen essential science advances’ for reforestation (Marshall et al., 2023) were recently
37 proposed. However, the literature that outlines current forestation frontiers does not consider
38 its effects on rainfall (Sheil et al., 2019).

39

40 It has long been recognized that forests participate in the regional hydrological cycle (see
41 discussions in Te Wierik et al., 2021; Bonan, 2023). Especially the negative effects of
42 deforestation in the Amazon on rainfall were noted already by some early influential papers
43 (Salati et al., 1979; Nobre et al., 1991; Shukla et al., 1990; Nepstad et al., 1994) and
44 corroborated on relatively local (Smith et al., 2023) to more regional scales (Staal, Flores, et
45 al., 2020). Trees pump water around from their roots via the stomata in their leaves to the
46 atmosphere, enhancing rainfall regionally (Spracklen et al., 2018). Therefore, deforestation,
47 especially but not exclusively in tropical areas such as the Amazon, may synergize with
48 climate-change-induced drought. Also in drier areas such as the Sahel, the importance of land
49 cover-rainfall feedback has been understood for decades (Savenije, 1995). It should, then, be
50 no surprise that the idea that ecological restoration could be used for the benefit of the
51 hydroclimate has been proposed along the margins of the wider restoration wave. Indeed, the
52 idea appears in notions as ‘conservation of water cycle on land via restoration’ (Makarieva et
53 al., 2006), ‘[water] supply-side thinking’ (Ellison et al., 2012), ‘aerial river management’ by
54 ‘smart reforestation’ (Weng et al., 2018), ‘induced precipitation recycling’ (Layton & Ellison,
55 2016), ‘moisture recycling governance’ (Keys et al., 2017), ‘moisture recycling as an
56 ecosystem service’ (Keys et al., 2016), ‘managing moisture recycling for nature-based
57 resilience’ (Ellison et al., 2019), ‘forest restoration for drought mitigation’ (Tuinenburg et al.,
58 2022), and ‘planting trees to combat drought’ (Baker, 2021). Across the globe, climate change
59 is creating or intensifying problems related to drought: forest fires are becoming more frequent
60 and intense, rivers are running dry, heat waves are getting more severe, and crop yields are
61 suffering from lack of water (IPCC, 2022a). It becomes clear that ecosystem restoration
62 projects may not only mitigate climate change by enhancing carbon sequestration but also
63 provide hydroclimatic buffers to already ongoing change.

64

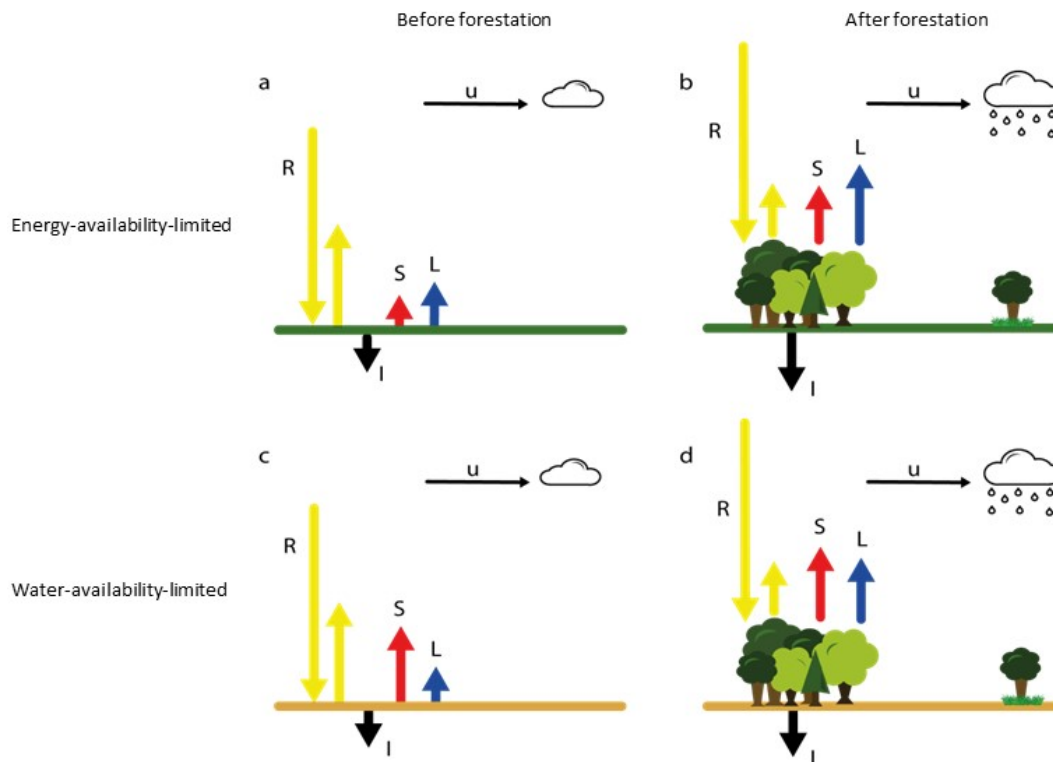
65 Building on these ideas and the latest published results, we here put forward that forestation
66 could be used to mitigate climate-change induced drying if it is done wisely (e.g. Hoek van
67 Dijke et al., 2022; Tuinenburg et al., 2022), but acknowledge that it may have adverse local

68 hydrological effects if done unwisely (see e.g. Filoso et al., 2017; Hoek van Dijke et al., 2022).
69 Specifically, we go one step further than the existing literature by arguing that—in addition to
70 other considerations such as biodiversity effects—it is starting to become possible to purposely
71 account for spatially explicit non-local rainfall effects of forestation in holistic reforestation
72 prioritization and climate-change vulnerability assessments globally. We call this concept
73 *targeted rainfall enhancement* (TRE). We explain that, despite a solid biophysical
74 underpinning and although results are now becoming available that allow for the first steps
75 towards TRE, more research is needed before reliable planning of TRE can be made that are
76 robust to projected climatic changes. We further suggest that moisture recycling networks
77 (Keune & Miralles, 2019; Wunderling et al., 2020; Wunderling, Wolf, et al., 2022; Zemp,
78 Wiedermann, et al., 2014) could be constructed and used to guide this TRE. In addition to
79 outlining the state-of-the-art regarding targeted rainfall enhancement through forestation, we
80 discuss how it can be viewed in the wider context of geoengineering. Lastly, we provide
81 avenues for future research in this Decade on Ecosystem Restoration.

82

83 **Biophysical underpinning of targeted rainfall enhancement**

84 Forests connect the global hydrological and energy cycles (Ellison et al., 2017). They promote
85 rainfall elsewhere by increasing atmospheric moisture content through enhanced
86 evapotranspiration (Keys et al., 2016). Therefore, the regional-scale impacts on rainfall, in
87 particular, depend on the evapotranspiration-enhancing properties of forests relative to other,
88 short-vegetation, types. Evapotranspiration can be considered forest-limited in areas where it
89 would be enhanced by increased forest cover. However, we can differentiate two ways in which
90 forests may lift the limitation for evapotranspiration by changing the local energy balance (Fig.
91 1). Firstly, under energy-availability-limited conditions, forestation darkens the land surface
92 and thus reduces albedo. This increases the absorbed solar radiation, making more energy
93 available for evapotranspiration (Bonan, 2008). Secondly, under water-availability-limited
94 conditions, forests are able to sustain higher rates of evapotranspiration by deeper roots that
95 allow them to store and access deep soil- and groundwater (e.g. Nepstad et al., 1994;
96 Seneviratne et al., 2010); by forest litter, large canopies, and branches and trunks hosting
97 epiphytes and lichens that allow for larger interception capacities (Savenije, 2018; Wang-
98 Erlandsson et al., 2014); as well as by more soil organic matter that promotes the infiltration
99 of rainwater into the soil (Y. Zhang et al., 2015).



100
 101 Figure 1: Simplified illustration of how forestation may enhance evapotranspiration through
 102 changing the energy balance. This may occur under both (a, b) energy-availability-limited and
 103 (c, d) water-availability-limited conditions. In energy-availability-limited systems, reduced
 104 albedo due to forestation leads to greater absorption of solar radiation R , increasing the amount
 105 of energy available for evapotranspiration, or latent heat (L), but also increasing sensible heat
 106 (S). Independent of how the additional energy is partitioned between latent and sensible heat,
 107 the moisture flux towards the atmosphere is expected to increase. In water-availability-limited
 108 systems, forestation increases not only access to deeper water layers but may also increase
 109 water infiltration (I) into the soil itself, alleviating the water limitation for evapotranspiration.
 110 In both types of systems, the increase in evapotranspiration may enhance rainfall downwind. u
 111 represents the horizontal wind flow that carries the moisture downwind to the target location
 112 of the enhanced evapotranspiration.

113
 114 In addition to (non-local) targeted rainfall enhancement, forestation may also be employed to
 115 enhance local rainfall (Theeuwens et al., 2023) (not shown in Fig. 1), but the net local
 116 hydrological effects of forestation are not always straightforward. Trees extract water from the
 117 landscape, but could also enhance rainfall by feeding convection, following an increase in the
 118 sensible heat flux (Branch & Wulfmeyer, 2019). However, especially in the tropics, forestation
 119 may cause cooling rather than warming, leading to greater atmospheric stability and hampering

120 upward air motion. Yet, it has also been proposed that a larger latent heat flux would locally
121 enhance rainfall if atmospheric moisture content is high enough. Close to saturation, a small
122 increase in atmospheric moisture would trigger rainfall, which would induce moisture
123 convergence and thus rainfall again (Makarieva et al., 2023) (not shown in the simple Fig. 1).
124 This would add to another biological pathway of local rainfall initiation via emissions of
125 biogenic volatile organic compounds that stimulate cloud formation (Peñuelas & Staudt, 2010).
126 However, empirical evidence indicates that, depending on spatial patterns, patchy deforestation
127 in tropical forests can locally also enhance rainfall through increasing turbulence (Lawrence &
128 Vandecar, 2015), implying that forestation that smoothens out such heterogeneous surface
129 roughness would locally reduce rainfall.

130

131 In summary, forest can enhance rainfall via different mechanisms, suggesting that it is possible
132 to target rainfall under a range of geographical and climatic conditions. However, to realize
133 spatially explicit assessments of targeted rainfall enhancement, we need more specific tools
134 than a mechanistic understanding of the processes only.

135

136 **Spatially resolving forest effects on rainfall**

137 Targeted rainfall enhancement relies not only on the extent to which enhanced
138 evapotranspiration reaches the desired target locations, but crucially also on how precise we
139 are able to identify these target locations. How much of the evapotranspiration from a certain
140 area ends up as rainfall on terrestrial systems depends on a range of meteorological conditions
141 that affect wind patterns and moisture recycling ratios. Although such regional effect of forests
142 on rainfall have been long recognized (see Te Wierik et al., 2021), only recently have there
143 been significant advances in atmospheric moisture tracking to understand atmospheric
144 moisture recycling on relatively fine spatial scales (down to 0.25° , around 25×25 km around
145 the equator) and the relatively fine temporal scale of a single month. These advances allow for
146 understanding rainfall effects of reforestation beyond large-scale, general estimates.
147 Specifically, the spatially explicit, regional connections between evapotranspiration and
148 rainfall can be reconstructed by the latest generation of atmospheric moisture tracking models
149 (Van der Ent et al., 2014; Tuinenburg & Staal, 2020; Keune et al., 2022). Although different
150 models (with different assumptions, complexity, and input data) exist (Van der Ent et al., 2013;
151 Tuinenburg & Staal, 2020), the same set of principles is shared among these moisture tracking
152 models: by using reanalysis data of evapotranspiration, atmospheric moisture profiles, wind
153 speeds and directions, and precipitation, moisture is ‘tracked’ through the atmosphere either

154 forward from evapotranspiration to precipitation or backward from precipitation to
155 evapotranspiration. A similar tracking approach has been implemented in the Weather
156 Research and Forecasting (WRF) model (Dominguez et al., 2016, 2022; Y. Gao et al., 2020;
157 Insua-Costa et al., 2022; Insua-Costa & Miguez-Macho, 2018; Yang & Dominguez, 2019).
158 This meteorological model is more complex than the moisture tracking models that are fed by
159 reanalysis data only, allowing users to experiment with different meteorological or land-cover
160 conditions, but also suffering from uncertainties inherent in the complexity involved. Earth
161 system models can also be used to study land-cover change effects on climate. However, not
162 many have specifically looked at forestation effects on rainfall and if so, these studies tend to
163 involve simultaneous large-scale forest cover changes, at regional-to-continental (e.g. Laguë
164 & Swann, 2016) or global scales (e.g. Portmann et al., 2022). They show that such massive
165 changes may affect atmospheric circulations at global scales and are thus useful to assess
166 effects of such rather extreme changes on these circulations. Yet, these models may be less
167 suitable to isolate detailed targeted rainfall enhancements due to their computational
168 requirements. Methods do exist, however, to differentiate between local (within a grid cell) and
169 non-local (from outside the grid cell) climatic effects of land cover change (Winckler et al.,
170 2017).

171
172 As with all models, to work with them requires experience and time investment, and especially
173 the latest models can be particularly data demanding. However, several global datasets of
174 evapotranspiration-to-precipitation links have recently been published, which can be used more
175 easily than the source code itself. The highest spatial resolution is achieved by Tuinenburg et
176 al. (2020), who published a dataset of mean-monthly climatologically averaged (2008–2017)
177 global evapotranspiration-precipitation connections between each pair of grid cells at 0.5°
178 resolution (~50×50 km around the equator). For these results, the Lagrangian model UTrack is
179 forced with the most detailed and latest (ERA5) hourly atmospheric reanalysis data for 25
180 atmospheric layers (Tuinenburg & Staal, 2020). Time series, however, of the destinations of
181 terrestrial evapotranspiration are published by Link et al. (2020), using the Eulerian moisture
182 tracking model WAM-2layers (Van der Ent et al., 2014) at 1.5° resolution.

183
184 In addition to spatially resolving moisture flows, understanding how land-cover changes affect
185 rainfall requires reliable estimates of the effects of those changes on evapotranspiration.
186 Different methods and approaches have been employed in order to estimate them. Spracklen et
187 al. (2012) used the back-tracked moisture flows themselves to do this: rainfall across the tropics

188 was related to the cumulative leaf area index (LAI) of the land surface along the trajectories
189 that the precipitated moisture had passed over. They found a positive correlation between
190 rainfall and the LAI along the trajectory, providing correlational evidence of the effects of
191 vegetation on rainfall. Meier et al. (2021) and Smith et al. (2023) used statistical approaches to
192 estimate local effects of forest change on rainfall, complemented in Meier et al. (2021) by a
193 trajectory-based approach to estimate remote effects. An alternative approach is to use
194 ecohydrological models in which actual evapotranspiration is calculated based on soil moisture
195 and vegetation properties (e.g. Wang-Erlandsson et al., 2014, 2018; Staal et al., 2018).

196

197 Moisture tracking has also been employed to study more local-scale moisture recycling. It was
198 found that, globally, the proportion of evapotranspiration that rains out is rather low: within
199 ~50 km it is in the order of 1–5% (Theeuwes et al., 2023). Nevertheless, in recycling models,
200 the process of fast recycling (local showers yielding rain before locally evaporated moisture is
201 fully mixed in the atmospheric column (Burde, 2006)) may generally be underestimated due to
202 the assumption of complete vertical mixing of locally evaporated moisture with advected
203 moisture (Dominguez et al., 2020).

204

205 In summary, atmospheric moisture tracking is a tool that can be used to move from a general
206 understanding for forest-atmosphere interactions to spatially and temporally resolving effects
207 of forestation. However, they cannot account for atmospheric circulation changes following
208 forestation, for which Earth system models are currently more suitable. Moisture tracking
209 within Earth-system and WRF models may offer better opportunities to study effects of specific
210 mechanisms on atmospheric moisture flows, whereas existing moisture flow databases offer
211 more user-friendliness. Intercomparison among different type of moisture tracking modelling
212 approaches and comparison with observation-based isotope studies are necessary to narrow the
213 uncertainty span.

214

215 **Benefits and trade-offs of targeted rainfall enhancement**

216 Although the above advances in atmospheric moisture tracking and other scientific progress
217 suggest that targeted rainfall enhancement (TRE; Figure 2) can be applied in principle, this
218 idea is further substantiated by recent work specifically on the effects of forest increase on
219 rainfall. Two independent studies recently assessed hydrological effects of the ‘global tree
220 restoration potential’ as estimated by Bastin et al. (2019), which amounts to an increase of 900
221 million ha of forest compared to present-day forest cover. The studies, by Tuinenburg et al.

222 (2022) and Hoek van Dijke et al. (2022), used the same moisture tracking model (UTrack), but
223 different methods to estimate the effects of forest cover change on evapotranspiration. Their
224 estimates of the globally averaged rainfall increase range between 4.8 mm yr^{-1} (Hoek van Dijke
225 et al., 2022) and 7 mm yr^{-1} (Tuinenburg et al., 2022). Two-thirds of the additional
226 evapotranspiration would precipitate over land; for 21%, this would be in areas that are
227 projected to become drier due to global climate change (Tuinenburg et al., 2022). However,
228 higher evapotranspiration is not always desired, as it can lead to lower water yield (precipitation
229 minus evapotranspiration, equivalent to streamflow) locally (e.g. Farley et al., 2005; Hoek van
230 Dijke et al., 2022). Indeed, in about half of the global land area, water yield would decrease; in
231 the other half, water yield would increase due to enhanced atmospheric moisture recycling
232 (Hoek van Dijke et al., 2022). This is consistent with the finding that the global increase in
233 evapotranspiration under high forestation would be greater than that in terrestrial precipitation
234 (Tuinenburg et al., 2022). A third study, by Cui et al. (2022), analyzed recent changes in global
235 leaf area and used output of the WAM2-layers model (Van der Ent et al., 2014; Link et al.,
236 2020) to assess how water yield has changed as a consequence of the leaf-area changes. They
237 found that as a global average, recent greening has led to an annual increase in water yield of
238 0.26 mm yr^{-1} . Importantly, they find that vegetation increase enhances water yield both locally
239 and in downwind areas in almost half (45%) of the globe. However, differences in rainfall
240 partitioning to infiltration, interception evaporation, and runoff could affect water yield (X.
241 Zhang et al., 2022). The above further illustrates that the hydrological effects of forestation are
242 not straightforward (Ellison et al., 2012) and highlight the importance of location (Staal, 2022):
243 whether effects of forestation on water yield are positive, positive but unacceptable (in case of
244 increased floods), negative, or negative but acceptable (if water is not limiting), depends on the
245 location of both the source of the moisture and its target.

246
247 Regardless of their hydrological effects, massive forest increases should not come at the
248 expense of native well-functioning ecosystems such as natural grasslands (Veldman et al.,
249 2015a, 2015b; Dudley et al., 2020). By excluding natural grasslands from the potential
250 forestation map, but also focusing on rainfall locations that are projected to become drier due
251 to global climate change, some preliminary focus areas for large-scale forestation-induced
252 targeted rainfall enhancement were already identified by Tuinenburg et al. (2022): the southern
253 and western Amazon, Mexico, eastern China, and Mediterranean Europe. However, depletion
254 of local water resources should also be prevented. Hoek van Dijke et al. (2022) name the
255 following regions where global forestation would enhance water yield: parts of the Sahel,

256 eastern Europe, southern Africa and South America, the Sahara, and the Himalayas, all of
257 which experience water scarcity for at least three months per year. Cui et al. (2022) attribute
258 the net positive effect of global greening on water yield mainly to recent greening in Europe,
259 western Siberia, parts of Africa and eastern China. These selections do not preclude the
260 potential of TRE to combat drying in other places, also due to caveats related to the methods
261 behind the global forest potential map (e.g. Veldman et al., 2019) and the scale of analysis.
262 Locally, restoration of forest patches that can sustainably provide relatively large amounts of
263 moisture to the atmosphere (e.g. gallery forests) could be prioritized. At the same time, critical
264 equity issues with regard to social considerations (e.g. on distribution, procedure, recognition,
265 or context of the ecosystem restoration) can decisively influence the success and effectiveness
266 of an ecosystem restoration project (Elias et al., 2022; Pascual et al., 2014). Further, a
267 disproportionately large number of people live in regions with a low Human Development Index
268 as well as high ecosystem restoration priority (Löfqvist et al., 2023). Therefore, an equity-
269 centered dimension is key for future ecosystem restoration projects, also with regard to TRE.
270 Which land-cover type gets replaced also matters for the TRE potential (Sterling et al., 2013).
271 Still, where they can be sustained, forests tend to provide the most stable flux of moisture to
272 the atmosphere compared to all other ones, as they can access deeper groundwater even during
273 periods of little to no rainfall. Hence, they buffer against droughts (Staal et al., 2018; O'Connor,
274 Dekker, et al., 2021) and may shorten dry seasons, such as in the Amazon (O'Connor, Santos,
275 et al., 2021; Wright et al., 2017).

276
277 Most of these results regarding active forestation so far remain hypothetical. However, one
278 notable example offers an empirical case study for large-scale ecosystem restoration, where
279 recent findings appear to provide a proof-of-principle, but also illustrate the complexities
280 involved. In the Chinese Loess Plateau, grassland area has increased by ~12,000 km² and forest
281 area by ~9000 km² between 1985–2015, largely by replacing decreasing croplands and barren
282 lands (Sun et al., 2022). As expected, evapotranspiration rates have increased in the region
283 (Feng et al., 2016; Shao et al., 2019; Tian et al., 2022). At the same time, tree cover increase
284 has improved soil water holding capacity and soil water availability in the area (Y. Zhang et
285 al., 2021), and importantly, for over 80% of the Loess Plateau, local-scale moisture loss due to
286 increased evapotranspiration has been compensated or surpassed by an increase in rainfall (B.
287 Zhang et al., 2022). However, in the remaining areas (primarily in arid regions, with mean
288 annual rainfall below 400 mm), water yield decreased. In these cases, forests were the main
289 revegetation type (B. Zhang et al., 2022). On average, rainfall has increased by 54 mm yr⁻¹,

290 due to both increased internal recycling and enhanced net moisture inflow into the Loess
291 Plateau, compared to an average increase in evapotranspiration of 23 mm yr⁻¹ (Tian et al.,
292 2022). It should, however, be noted that a simultaneous reduction in agricultural water use may
293 obscure some of the actual water loss caused by the forestation efforts (Zhou et al., 2020). The
294 case of the Loess Plateau illustrates the importance of targeted rainfall enhancement to promote
295 rainfall, while avoiding the negative effects of forestation on water storage (Zhao et al., 2021),
296 and on streamflows and downstream wetlands (Xi et al., 2022).

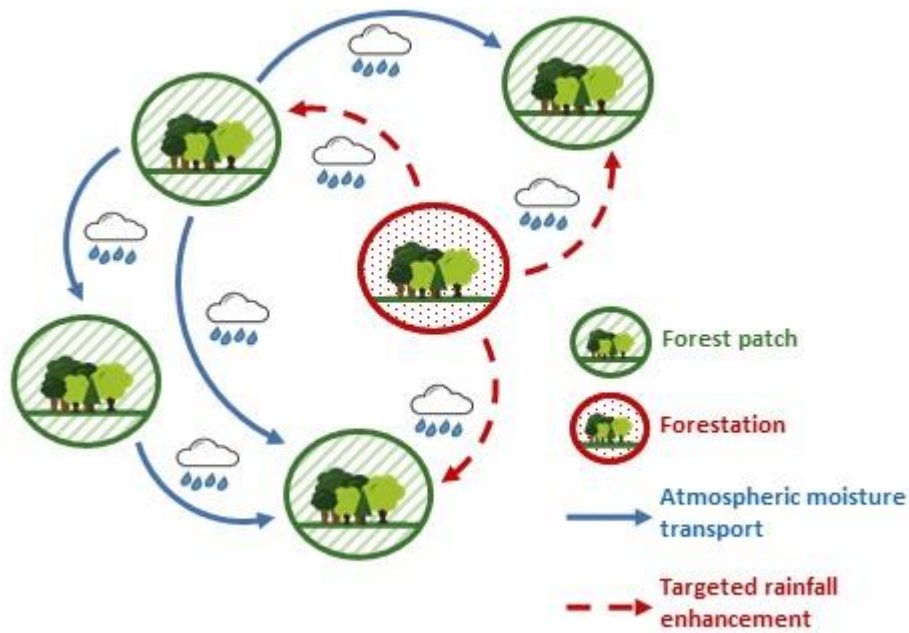
297

298 Many benefits of forestation, such as carbon capture, will occur on the time scale of decades.
299 TRE, however, is an ecosystem service that will return relatively fast after forestation. As soon
300 as photosynthesis is restored, given sufficient rooting depth, evapotranspiration will increase,
301 although in measurements from Japan, evapotranspiration continued to increase for 20 years,
302 after which it seemed to stabilize (Murakami et al., 2000). Measurements from France suggest
303 that, although the partitioning of evapotranspiration depends on stand age, total
304 evapotranspiration is independent of stand age (Delzon & Loustau, 2005). Forest recovery is
305 fastest and easiest in wetter regions such as in the tropics (Poorter et al., 2016), but those areas
306 may not be the regions that would benefit most from enhanced atmospheric moisture recycling.
307 Thus, there might be a trade-off between ease of recovery and rainfall benefits of recovery.

308

309 In summary, forestation should not be seen as a silver bullet for solving water shortages either
310 locally or regionally, so each project should be assessed individually to prevent adverse
311 hydrological effects. However, increasing evidence indicates that targeted rainfall
312 enhancement through forestation would be beneficial in some cases.

313



314

315 Figure 2: Targeted rainfall enhancement as an objective of forestation. Forestation tends to
 316 enhance atmospheric moisture transport, which would increase rainfall in target areas directly,
 317 but may also have cascading effects in the forest system. These cascading effects may feed
 318 back to the original forested patch, adding to other local effects of forestation.

319

320 **Is targeted rainfall enhancement geoengineering?**

321 It has not escaped our notion that if either afforestation or reforestation becomes more
 322 deliberate and planned with regard to its (hydro)climatic effects, it may be shifting from the
 323 domain of ecosystem restoration to geoengineering, which tends to be perceived differently
 324 and raise different ethical concerns. The IPCC defines both afforestation and reforestation as
 325 mitigation, that is, a human intervention to reduce the sources or enhance the sinks of
 326 greenhouse gases, whereas geoengineering is seen as efforts to stabilize the climate system by
 327 directly managing the energy balance of the Earth, thereby overcoming the enhanced
 328 greenhouse effect (Minx et al., 2018). Tree planting does, however, affect the energy balance
 329 of the Earth and could thus affect regional rainfall levels similar to marine cloud brightening,
 330 which is widely considered to be geoengineering (Jones et al., 2009; Latham et al., 2012). TRE
 331 could also be interpreted as geoengineering when defined as an ‘intentional large-scale
 332 manipulation of the environment, particularly manipulation that is intended to reduce undesired
 333 anthropogenic climate change’ (Keith, 2000). Given those definitions we think that some
 334 reflections within this context are warranted.

335

336 TRE indeed entails reducing undesired anthropogenic climate change through manipulation of
337 the environment; the scale from which it can be considered geoengineering, however, is
338 unclear. Planting a tree for environmental benefits, for instance shading, is a form of managing
339 the local climate and in IPCC terms, adaptation. At larger scales, with the intention of realizing
340 negative carbon emissions, this turns into mitigation. However, the biophysical effects of full
341 forest potential are likely to have such large global effects on the radiation balance and moisture
342 recycling (Portmann et al., 2022) that realizing it could be considered geoengineering. Despite
343 nonlinearity in the effects of forest expansion on rainfall (Baudena et al., 2021; Lawrence &
344 Vandecar, 2015; Zemp, Schleussner, et al., 2014), there is no clear transition from small-scale
345 non-geoengineering measures to large-scale geoengineering. In principle, if forestation is
346 strictly *reforestation* in the sense that previous deforestation is reversed, then it can reasonably
347 be considered ecosystem restoration even if TRE is taken into account.

348

349 Simply *reforestation*, considering historical baseline conditions, may, however, not suffice. A
350 Holocene baseline of forest distributions (Steffen et al., 2015) and the hydrological cycle
351 (Wang-Erlandsson et al., 2022) can be justified as a safe space to operate in, meaning that it
352 can serve as a guide to prioritize restoration without it pushing measures outside the historical
353 domain. As planting trees where they did not occur before would coincide with larger
354 unknowns, the term ‘geoengineering’ may apply. However, a key question is whether we can
355 always assume no harm from reforestation. Locally, Holocene conditions were not always
356 constant (e.g. Dermody et al., 2012) and planting forests may not reverse past anthropogenic
357 effects but add to them (Heck et al., 2016). Furthermore, as discussed, different climate futures
358 may imply different reforestation strategies: restoring forests that fed rainfall to certain areas
359 in the past may not do so in the future. Holocene conditions at the global scale may still imply
360 major redistribution locally and regionally. Thus, it can be argued that controlling forest
361 distributions for those reasons, even when reforestation in a strict sense is employed, is
362 geoengineering.

363

364 Beyond definition questions of geoengineering, TRE involves ethical considerations around
365 restoration in general. Redistributions of biomass and hydroclimatic conditions are prone to
366 have both winners and losers, and ensuring justice can be necessary for gaining social and
367 political acceptance. Already, all kinds of restoration efforts consider multi-functional use of
368 grasslands for biodiversity restoration, or multiple-use of forests for climate mitigation. Such

369 assessments are not neutral and necessitates an ethically explicit judgement (Batavia & Nelson,
370 2018). Further, it can be argued that geoengineering through plantation is already happening in
371 many ways. For instance, a shift is currently happening from planting of seedlings grown from
372 locally sourced seeds (geographically based reforestation) to genomics-based assisted
373 migration by selecting seeds based on expected future climates (Findlater et al., 2022). Forest
374 managers also increase climate benefits by using more reflective and deciduous tree species
375 (Jackson et al., 2008). Those combined effects of carbon sequestration and biophysical effects
376 in terms of energy and hydrology are not taken into account in current climate policy (Boucher
377 et al., 2014). Thus, a debate around TRE would be in line with different forest management
378 efforts already taking place and helps ensure that justice issues related to moisture
379 redistribution are accounted for.

380

381 **Outlook**

382 What could be considered sensible targeted rainfall enhancement currently, may not be so in
383 the future. Firstly, global climate change may affect global atmospheric moisture flows and
384 thereby regional moisture connections (Baker & Spracklen, 2022). Other relevant hydrological
385 variables will probably also change: the timing and intensity of evapotranspiration and rainfall
386 (IPCC, 2021); water-use efficiency following CO₂ fertilization (Dekker et al., 2016); and the
387 moisture holding capacity of the atmosphere, resulting in different recycling rates (Dominguez
388 et al., 2006) and distances (Gimeno et al., 2021). The continued drying and warming trends can
389 also lead to changes in prioritization of forestation areas to stop remote carbon emissions (Staal
390 et al., 2023; Pires, 2023) and to shifting forest suitability ranges (Staal, Fetzer, et al., 2020).
391 Secondly, TRE should be robust under land-cover changes, which includes TRE itself. For
392 instance, large-scale forest cover changes may alter temperature and pressure gradients and,
393 consequently, moisture flows. This may cause certain other areas to either become more or less
394 suitable for TRE such that global or regional TRE potential feeds back to itself (Tuinenburg et
395 al., 2022). Therefore, to better account for the complex and dynamic nature of the Earth system,
396 we need explicit integration of future simulated evapotranspiration and rainfall flows with
397 harmonized land-use change scenarios including better integration with ecohydrological
398 processes, as well as wind patterns and atmospheric moisture content at high spatial and
399 temporal resolution for a range of scenarios. The ‘tagging’ of moisture in an Earth system
400 model (Harrington et al., 2023) would be a logical and significant step forward in this regard.

401

402 A network perspective on targeted rainfall enhancement could inform about possible target
403 locations to strategically enhance moisture flows (Figure 2). In previous work, atmospheric
404 moisture connections among forest patches (in the order of 100 km) have been converted to
405 directed networks, here meaning one-way flows between moisture sources and sinks (Zemp,
406 Wiedermann, et al., 2014; Krönke et al., 2020). By analyzing the networks' structures for their
407 role in overall dynamic behavior (Wunderling et al., 2020; Wunderling, Wolf, et al., 2022), it
408 is possible to, for instance, identify major transmitters of moisture flows. Thus, the
409 'importance' of a forest patch on the entire forest-rainfall network can be assessed. Forestation
410 may significantly enhance evapotranspiration and can thus add particular evaporation-to-
411 rainfall links to the network. It may also enhance transmission of moisture previously
412 contributed by forest to the network. As such, counterintuitively, forestation could enhance the
413 importance of upwind forest patches to the overall forest system. Interestingly, it is consistently
414 found in the network literature that networks are usually very robust to random removal of
415 links, but considerably less robust to targeted attacks and removals of links (Albert et al., 2000;
416 Newman, 2018)—a 'targeted attack' in this context meaning deliberate restoration of important
417 patches and 'targeted removal' meaning deliberate deforestation of important patches. Indeed,
418 deforestation of particular patches in the Amazon may lead to disproportional loss of resilience
419 of the full forest-climate system (Wunderling, Staal, et al., 2022). Targeted rainfall
420 enhancement, in turn, could then also disproportionately enhance forest rainfall by stabilizing
421 the network as a whole (Figure 2), as well as connecting different parts of the network that
422 were not interacting before (Chen et al., 2007). By tapping into the numerous ways to determine
423 and tune the resilience of networks (e.g. in network control theory; Liu et al., 2011; J. Gao et
424 al., 2016), explicit assessments of moisture-recycling network structures and dynamics
425 (Wunderling et al., 2021; Wunderling, Wolf, et al., 2022) could become an important tool for
426 realizing targeted rainfall enhancement.

427

428 Despite an emerging scientific basis for TRE, major uncertainties remain. The foundations for
429 our arguments rest on model simulations fed by inputs from observation-based data, and major
430 model developments are still required for reliable assessments. Refined estimates that account
431 for cross-scale dynamics require coupled modelling, including vegetation-induced changes in
432 atmospheric circulation (Portmann et al., 2022) and climate-change-induced modifications of
433 moisture flows (Findell et al., 2019). Care should also be taken to fit forestation efforts in line
434 with prevailing and projected future hydroclimate. Importantly, TRE should not be a stand-
435 alone consideration, but integrated into existing or emerging frameworks for restoration and

436 conservation prioritization (Brancalion et al., 2019; Griscom et al., 2020; Strassburg et al.,
437 2020; Jung et al., 2021; Hua et al., 2022; Aronson et al., 2020). This includes accounting for
438 water-related trade-offs and co-benefits such as changes in water yield and groundwater
439 recharge, as well as broader sustainability and ethical concerns. Through a balanced inclusion
440 of TRE in forest ecosystem service provisioning valuations, it could also add to the value of
441 these services (Asbjornsen et al., 2022). As such, TRE can become part of holistic assessments
442 of reforestation potential in the current Decade on Ecosystem Restoration and beyond.

443

444 **Conclusion**

445 A number of recent papers have discussed or hinted at the possibility of using forests to manage
446 rainfall. Considered together, it emerges that the effects on rainfall of potential forestation are
447 significant and relevant. Forestation programs would mitigate not only global climate change
448 itself, but also its adverse effects in the form of drying. With increasingly better data and
449 models, we argue that targeted rainfall enhancement (TRE) now can be considered by and
450 integrated in forestation initiatives and conservation prioritization frameworks.

451

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463

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